www.the-innovation.org

Iron-rich Fe-O compounds at Earth's core pressures

Jin Liu,^{1,2,9,*} Yang Sun,^{3,9} Chaojia Lv,¹ Feng Zhang,⁴ Suyu Fu,⁵ Vitali B. Prakapenka,⁶ Caizhuang Wang,⁴ Kaiming Ho,⁴ Jungfu Lin,^{5,*} and Renata M. Wentzcovitch^{3,7,8,*}

*Correspondence: jin.liu@hpstar.ac.cn (J.L.); afu@jsg.utexas.edu (J.L.); rmw2150@columbia.edu (R.M.W.)

Received: July 22, 2022; Accepted: November 13, 2022; Published Online: November 15, 2022; https://doi.org/10.1016/j.xinn.2022.100354 © 2022 The Authors. This is an open access article under the CC BY-NC-ND license (http://creativecommons.org/licenses/by-nc-nd/4.0/).

GRAPHICAL ABSTRACT



PUBLIC SUMMARY

- A family of Fe-rich Fe_nO compounds is revealed under Earth's core conditions
- The FenO compounds feature close-packed Fe and O monolayers
- The separate O monolayer stabilizes FenO compounds at multi-megabar pressures
- The Earth's inner core can contain more oxygen than expected due to FenO phases
- It alters our understanding of thermal evolution and seismic features of the core



Iron-rich Fe-O compounds at Earth's core pressures

Jin Liu,^{1,2,9,*} Yang Sun,^{3,9} Chaojia Lv,¹ Feng Zhang,⁴ Suyu Fu,⁵ Vitali B. Prakapenka,⁶ Caizhuang Wang,⁴ Kaiming Ho,⁴ Jungfu Lin,^{5,*} and Renata M. Wentzcovitch^{3,7,8,*}

¹Center for High Pressure Science and Technology Advanced Research, Beijing 100094, China

²CAS Center for Excellence in Deep Earth Science, Guangzhou 510640, China

³Department of Applied Physics and Applied Mathematics, Columbia University, New York, NY 10027, USA

⁴Department of Physics, Iowa State University, Ames, IA 50011, USA

⁵Department of Geological Sciences, Jackson School of Geosciences, The University of Texas at Austin, Austin, TX 78712, USA

⁶Center for Advanced Radiation Sources, University of Chicago, Chicago, IL 60439, USA

⁷Department of Earth and Environmental Sciences, Columbia University, New York, NY 10027, USA

⁸Lamont–Doherty Earth Observatory, Columbia University, Palisades, NY 10964, USA

⁹These authors contributed equally

*Correspondence: jin.liu@hpstar.ac.cn (J.L.); afu@jsg.utexas.edu (J.L.); rmw2150@columbia.edu (R.M.W.)

Received: July 22, 2022; Accepted: November 13, 2022; Published Online: November 15, 2022; https://doi.org/10.1016/j.xinn.2022.100354

© 2022 The Authors. This is an open access article under the CC BY-NC-ND license (http://creativecommons.org/licenses/by-nc-nd/4.0/).

Citation: Liu J., Sun Y., Lv C., et al., (2023). Iron-rich Fe-O compounds at Earth's core pressures. The Innovation 4(1), 100354.

Oxygen and iron are the most abundant elements on Earth, and their compounds are key planet-forming components. While oxygen is pervasive in the mantle, its presence in the solid inner core is still debatable. Yet, this issue is critical to understanding the co-evolution and the geomagnetic field generation. Thus far, iron monoxide (FeO) is the only known stoichiometric compound in the Fe-FeO system, and the existence of iron-rich Fe_nO compounds has long been speculated. Here, we report that iron reacts with FeO and Fe2O3 at 220-260 GPa and 3000-3500 K in laser-heated diamond anvil cells. Ab initio structure searches using the adaptive genetic algorithm indicate that a series of stable stoichiometric FenO compounds (with n > 1) can be formed. Like ε -Fe and B8-FeO, Fe_nO compounds have close-packed layered structures featuring oxygen-only single layers separated by iron-only layers. Two solid-solution models with compositions close to Fe₂O, the most stable Fe-rich phase identified, explain the X-ray diffraction patterns of the experimental reaction products quenched to room temperature. These results suggest that Fe-rich FenO compounds with close-packed layered motifs might be stable under inner core conditions. Future studies of the elastic, rheological, and thermal transport properties of these more anisotropic Fe_nO solids should provide new insights into the seismic features of the inner core, inner core formation process and composition, and the thermal evolution of the planet.

INTRODUCTION

The composition and the nature of the solid phases in the Earth's core remain largely uncertain.¹ While iron is the major component in Earth's liquid outer core (OC) and solid inner core (IC), seismology requires light elements to account for the IC and OC's "density deficit" relative to pure iron.^{2,3} S, Si, C, H, and O are the most likely light elements in the liquid core.⁴ Light elements may be incorporated into the solid IC as the OC continuously solidifies by forming iron-rich compounds,^{1,5,6} such as Fe₇C₃, Fe₃S, and Fe₃Si. High-pressure melting experiments found that eutectic Fe-FeO liquids can contain much oxygen, e.g., ~15 wt % at IC boundary conditions.⁷ Modeling of core-mantle differentiation also indicates that oxygen is a vital light element in the liquid core.⁸ The Fe-O system has shown rich stoichiometry under oxygen-rich conditions, e.g., the recently discovered FeO₂ under lower-mantle conditions.^{9,10} However, oxygen has been excluded from the IC because of its limited solubility in solid iron observed at low pressures and the absence of known Fe-rich Fe-O compounds.^{11,12} The discovery of stable, solid Fe-O alloys at core pressures reported in this article changes this view, potentially impacting our understanding of the origin of the density deficit contrast between the IC and the OC and the OC composition, IC formation process, evolution, properties. etc.

Few studies have attempted to identify such O-bearing Fe-rich solids. In particular, *ab initio* studies by Sherman¹³ and Alfè et al.¹⁴ addressed substitutional and interstitial FenO structural models that were unstable against decomposition into the end members Fe and FeO. Weerasinghe et al.¹⁵ performed a crystal structure search on the Fe–O system and predicted the possible existence of low-enthalpy, iron-rich FenO phases under core pressures. Still, none have yet been synthesized at core pressures. Previous studies have shown that FeO transforms from an insulating inverse-B8 phase to a metal with the NiAs-type (B8) structure at pressures higher than 120 GPa.^{16,17} The latter features alternating single-species (Fe-only or O-only) close-packed layers with layered motifs similar to ϵ -Fe. Pure oxygen also becomes metallic at 96 GPa.¹⁸ The change of the chemical character of FeO and O at high pressures may change their solubility in solid iron at megabar pressures.¹⁹ Here, we combine high-pressure and -temperature (P-T) experiments and crystal structure predictions to search for iron-rich Fe-O compounds at pressures above 200 GPa. Because crystal structure searches identify low-temperature phases, we solve the X-ray diffraction (XRD) patterns of samples produced in P-T experiments but quenched to room temperature.

RESULTS AND DISCUSSION Experimental synthesis

We performed five high-P-T experiments to synthesize FenO compounds. We carried out synchrotron XRD experiments on the Fe-FeO system under high-P-T conditions using laser-heated diamond anvil cells at the Advanced Photon Source at the Argonne National Laboratory and the Shanghai Synchrotron Radiation Facility. The starting materials were mixtures of pure iron with FeO or Fe₂O₃ powder in a size range of submicron to a few microns (refer to details in the material and methods). Repetitive micro-second laser pulses were employed to heat the sample for a short period of ~1 s. This laser-heating strategy significantly reduced potential carbon diffusion from the diamond anvils into the sample chamber (see Note S1). The samples were well insulated from the diamond anvils by the KCl pressure-transmitting medium and were heated up evenly on both sides to reach high temperatures readily. No reaction was observed in the first three runs at 100-200 GPa up to 2600 K (Figure S1), consistent with previous experimental studies of FeO.^{12,20} In run 4, the Fe-FeO mixture was compressed firstly to 220 GPa at room temperature. Upon laser pulse heating to 2000–3000 K, diffraction peaks of both ϵ -Fe and B8-FeO merely became sharper, without the appearance of any new reflections. Further laser pulse heating to 3000-3200 K showed dramatic XRD pattern changes in a few seconds (left panel in Figure S2). The integrated diffraction peak intensities of the corresponding Fe and FeO phases reduced rapidly, indicating that iron reacted with FeO above 3000 K and 220-230 GPa. A similarly rapid chemical reaction happened in run five, starting from a $Fe-Fe_2O_3$ mixture that was compressed to 200–250 GPa at room temperature and exposed to repetitive laser pulse heating. After reaching \sim 3000 K, the peak intensities of the Fe and Fe₂O₃ phases decreased, while a set of additional diffraction peaks emerged (right panel in Figure S2). At least until 260 GPa and 3500 K, they remained unaltered before the anvils shattered upon further heating.

The diffraction peaks of the new phases did not change significantly after quenching the sample to room temperature (see Figures 1 and S2). The quenched new phase(s) remain (meta)stable at room temperature upon decompression to ~160 GPa. To improve the quality of the quenched samples'



Figure 1. Typical X-ray diffraction patterns of quenched samples collected before and after the reaction of Fe with FeO at 220 GPa and Fe₂O₃ at 245 GPa at high temperatures Left panel: B8-FeO has lattice parameters a = 2.3480(17) and c = 4.7901(41) Å; hcp–Fe has a = 2.1792(19) and c = 3.4865(32) Å; hcp-Re has a = 2.4733(23) and c = 3.9944(47) Å; cubic B2-KCl has a = 2.7292(53) Å at 215-220 GPa and room temperature; B2 denotes B2-FeO. Right panel: starting materials were a mixture of iron and hematite; hcp-Fe has a = 2.2195(81) and c =3.5220(115) Å at 190 GPa, and a = 2.1489(68) and c = 3.4391(95) Å at 245 GPa and room temperature; Fe₂O₃ has a = 2.2903(57), b = 7.3913(116), and c = 5.5172(81) Å at 190 GPa, and a = 2.2412(45), b = 7.2328(103), and c = 5.3989(81) Å at 245 GPa and room temperature; cubic B2-KCl has a = 2.7705(64) Å at 190 GPa and a = 2.7024(71) Å at 245 GPa at room temperature. Vertical ticks: ε-Fe (black), B2-FeO (red), B8-FeO (magenta), post-perovskite Fe₂O₃ (orange), B2-KCI (olive), and Re gasket (blue). Blue shaded areas highlight some of the new diffraction peaks that emerged at the expense of Fe and FeO/ Fe₂O₃ peaks when the temperature reached 3000-3200 K at 220 GPa and 3300-3500 K at 245 GPa, respectively (see Figures S3 and S4 for the indexing to the reaction products).

tent with the stacking in $\epsilon\text{-Fe}$. Layer sequences involving oxygen, e.g., Fe-O-Fe or Fe-Fe-O, favor ABC stacking over ABA. Five arrangements involving adjacent O-O layers, e.g., Fe-O-O or O-Fe-O with ABA

XRD patterns, we rotated the diamond anvil cell (DAC) by ±10° around the incident X-ray beam direction. This procedure significantly improved the XRD peak intensities relative to the background in these polycrystalline samples (Figures S3 and S4). The short 1-s data collection time makes this procedure very challenging at high *P*–*T*, and the high-temperature XRD patterns show a relatively higher background intensity (Figure S2). This was another reason for using the XRD patterns collected at room temperature and 220–245 GPa to identify the new phases. The new diffraction peaks from Fe + FeO and Fe + Fe₂O₃ reactions show consistent *d* spacings (Figure S5), indicating that similar phases are produced under different experimental conditions. The new peaks could not be indexed to high-pressure phases of known oxygen-rich Fe–O compounds (e.g., Fe₄O₅, Fe₅O₆, Fe₅O₇, Fe₂O_{3+δ}, and FeO₂ in Figure S6),^{10,21–24} or iron carbides (e.g., Fe₇C₃ and Fe₃C in Figure S7).^{25–27}

Crystal structure prediction

To solve the XRDs, we first conducted crystal structure searches for Fe_nO (n > 1) compounds using the adaptive genetic algorithm (AGA) at 215 GPa.²⁸ By searching in a wide range of Fe-rich compositions (see material and methods for details), several new stoichiometric Fe_nO phases are found to be more stable than the combination of two end members, ε -Fe and B8-FeO. Figure 2A shows the convex hull obtained from these crystal structure searches. Four compounds with Fe₃O (space group *P*6₃/*mmc*), Fe₂O (*P*3*m*1), Fe₃O₂ (*R*3*m*), and Fe₄O₃ (*R*3*m*) stoichiometries are ground states at 215 GPa. The Fe₆O₅ (*R*3*m*) phase is very close to the convex hull, with only 0.2 meV/atom above it. All these new Fe-rich compounds have single species (Fe-only or O-only) hexagonal close-packed layers, similar to the end-member phases ε -Fe (*P*6₃/*mmc*) and B8-FeO (*P*6₃/*mmc*) (Figure 2B). Furthermore, most low-enthalpy compounds also have hexagonal close-packed layers, as noted by the red symbol in Figure 2A.

To study the close-packing motifs in these phases, we selected structures whose formation enthalpies are within 26 meV/atom (\sim 300 K) above the convex hull. The lattice parameter *c* of these structures displays nearly linear dependence on the number of layers in the primitive cell (Figure 2C), suggesting a similar layered motif in all these structures. By examining these structures' stacking sequences, we found a simple compound formation rule. Only seven different three-layer stacking sequences out of 12 possible ones are present in these low-energy structures (see Figure 2D). The Fe–Fe–Fe layer sequence favors ABA stacking more than ABC, consisstacking are absent among these structures. Therefore, O-only layers are well separated from each other. This stacking rule explains why the previously proposed Fe_nO models are unstable,^{13,14} which included oxygen as substitutional and interstitial ions. The Bil₃-type (*R*-3) model is hexagonal,¹⁴ but the oxygen layers are not closely packed. All those atomic arrangements have high enthalpy. A previous computational search found an Fe₃O hexagonal structure ($P \ \overline{6} \ m2$)¹⁵ more similar to the *P*6₃/*mmc* ground state revealed here. However, the enthalpy of the *P* $\overline{6} \ m2$ phase is still ~10 meV/atom higher than the *P*6₃/*mmc* phase (Figure S8). Further calculations up to IC pressures confirm the stability of *P*6₃/*mmc* Fe₃O against the decomposition into Fe and FeO (Figure S8). The dynamic stability of these new phases is also confirmed at both 215 and 380 GPa (Figure S9). High temperatures typical of the IC might introduce anharmonic effects that should be investigated carefully using appropriate methods for high-temperature calculations.^{29–31}

XRD solutions

First, we compared the predicted XRD patterns of the computationally identified Fe_nO phases with the experimental ones. The four Fe_nO phases in the convex hull can partially match the XRD lines of the quenched reaction products but do not provide a complete interpretation of the diffraction patterns (Figure S10). By including a few metastable Fe_nO phases, the multi-phase aggregate XRD matches the experimental data better at both 220 and 245 GPa (Figure S11). Considering that all these Fe_nO phases follow the same basic stacking rule, we anticipate the entropic stabilization of disordered solid solutions at high temperatures. To clarify this point, we set up reverse Monte Carlo (RMC) simulations of supercell structures with up to 60 layers (~10 nm) to search for the Fe/O stacking sequence that best matches the experimental XRD patterns. The interlayer distances were set using the most probable spacings found in the crystal structures identified with the AGA (see Figure 3A and refer to material and methods for more RMC simulation details).

The XRD deviation, D_{XRD} , which is the criterion used to select preferred structures during the RMC sampling, is defined as $D_{XRD} = \sqrt{(I_{exp}(2\theta) - I_{sim}(2\theta))^2}$, i.e., the mean-square deviation between the simulated, I_{sim} , and the experimental, I_{exp} , peak intensities. The 2θ range analyzed is restricted to $9.2-10.2^\circ$ at 220 GPa and $10.6-11.6^\circ$ at 245 GPa, corresponding to the prominent peaks of the new reaction products



Figure 2. Crystal structure of identified Fe_nO compounds at 215 GPa (A) The formation enthalpy of AGA-searched compounds referenced by the dashed line shows their relative stability relative to the decomposition into ε -Fe and B8-FeO. The solid line indicates the convex hull formed by the thermodynamically stable compounds. Purple circles show structures with hexagonal close-packed layers. (B) Crystal structures of four Fe-rich compounds in the ground states. Gold and red solid symbols represent iron and oxygen, respectively. (C) Lattice parameter c of low-energy close-packed crystals as a function of the number of layers. (D) Frequency of structural and chemical order in the nearest neighbor layers of the low-energy close-packed crystals.

in Figure 1. Figure 3B shows the D_{XRD} values for the different supercell sizes and iron concentrations obtained in the RMC simulations. The smallest D_{XRD} at 220 GPa points to a 42-layer structure with chemical composition Fe₂₈O₁₄ shown in Figure 3C. This structure's XRD matches all additional experimental peaks of the reaction products. Ab initio calculations show the formation enthalpy of this phase is only 28 meV/ atom above the convex hull. This supercell's calculated equation of state agrees well with the experimental one (see Note S2 and Figure S12), confirming that this structure and composition are a reasonable solution for the experimental XRD. A similar RMC was performed to solve the XRD of the reaction products of Fe and Fe_2O_3 at 245 GPa (Figure S13). A 38-layer structure solution with composition Fe25O13 can match the experimental data well, as shown in Figure S3. Analyses of these structural models indicate that they represent structures lacking long-range stacking or chemical order. However, they display medium-range structural order, i.e., domains with the preferred stacking order on the scale of \sim 20 atomic layers (~30 Å) (see Note S3 and Figures S14-S16). Some short-range-order domains are also observed, e.g., the 5-layer sequence FeOFeOFe appears several times in Figures S15C, S15D, S16A, and S16B. The combination of and interference between the medium-range structural order domains give rise to the multi-peak patterns in the experimental XRD. The new reaction products likely consist of a disordered alloy with compositional fluctuation in the structure. The rapid temperature guench of these entropically stabilized high-temperature phases retains the disordered nature of the high-temperature phase. The high IC temperatures may favor the stabilization of these high-entropy phases. On the other hand, the long geological timescale may equilibrate an aggregate also containing the stoichiometric Fen0 phases predicted on the convex hull. The formation of a single disordered phase versus an aggregate containing the stoichiometric FenO compounds depends on the entropy versus enthalpy competi-

tion in FenO, which will define the high-P-T Fe-FeO phase diagram at relevant conditions.

Electronic density of states

Figure 4 shows the electronic density of states (DOS) of ε -Fe, Fe₃O, Fe₂O, and B8-FeO at 215 GPa and T_{el} = 3000 K. All Fe_nO phases discovered are metallic. In ε -Fe, the Fermi level falls on a valley (Figure 4A), while in B8-FeO, it falls on a peak. Consistently, in Fe_nO, the carrier density increases systematically with increasing oxygen content. In Fe₂O, the Fermi level falls on a shallow valley, and the DOS has features resembling those of ε -Fe and FeO. Fe₃O has two kinds of iron: those with (Fe1, Fe2, Fe4, Fe5) and those without (Fe3, Fe6) oxygen first neighbors (see Figure 4). Fe3-type iron's contribution to the DOS (red area in Figure 4B) resembles that of ε -Fe, while Fe1-type iron's contribution (blue area in Figure 4B) resembles Fe₂O's.

The chemical bonding nature in the Fe_nO phases is analyzed using electron localization function (ELF) and Bader charge transfer.^{32–34} Figure 5 shows the ELF cross-sections in the (110) plane containing both Fe and O atoms for ϵ -Fe, Fe₃O, Fe₂O, and B8-FeO. In ϵ -Fe, the ELF is relatively uniform. In B8-FeO, the valence electrons are more concentrated near oxygens, while Fe layers maintain a relatively uniform electron density. In Fe₃O and Fe₂O, adjacent Fe-Fe layers show similar ELFs to ε -Fe, while adjacent Fe-O layers have ELFs like those in B8-FeO. Table 1 shows Bader charge transfers in these compounds. Approximately 0.97 electrons transfer from Fe to O in B8–FeO. In Fe₃O and Fe₂O, O ions gain roughly the same number of electrons, but the electron loss per Fe in Fe₂O is 0.47, i.e., half of that in FeO. In Fe₃O, Fe without O neighbors (Fe3/Fe6) shows almost no electron loss, while Fe with O neighbors (Fe1/Fe2/Fe4/Fe5) shows an electron loss of 0.45, similar to Fe in Fe₂O. Therefore, the electronic DOS, chemical bonding, and charge transfer in the FenO phases can be interpreted by recalling those in the end-member phases, i.e., ϵ -Fe and B8-FeO. This seemingly simple and sort-of-short-range interaction between Fe and B8-FeO





suggests that several stoichiometries are possible in the Fe–FeO system, i.e., Fe_nO phases containing O-only and Fe-only close-packed layers stacked in different ways that depend on n. An entropically stabilized alloy with a stacking disorder is likely to form at high temperatures.

Geophysical implications

Our high-P-T experiments show that Fe and FeO/Fe2O3 react at 220-260 GPa at temperatures higher than 3000 K. At the same time, our ab initio crystal structure prediction identifies several stable Fe_nO phases with n > 1 at similar pressure conditions. These newly identified phases differ from previously suggested Fe_nO models^{13,14} and consist of single species (Fe-only or O-only) close-packed layers in various stacking sequences with an O-only single layer always separated by Feonly layers. The best solutions for the experimental XRD peaks of the new phases quenched to room temperature point to structures lacking long-range chemical or stacking order. However, they still have average compositions close to Fe₂O, i.e., that of the most stable phase identified computationally in the Fe-FeO system. With this in mind, the new high-temperature reaction products are likely entropically stabilized solid solutions whose intrinsic disorder is retained upon quenching. Notably, these high-entropy phases may be favored over stoichiometric compounds (e.g., Fe₂O and Fe₃O) at high IC temperatures. In general, the presence of oxygen layers should make Fe_nO alloys more anisotropic from an elastic and thermal transport perspective compared with pure iron. These properties must be carefully investigated since they are essential for understanding the core's thermal evolution and seismic features, including the IC's age, energy source powering the geodynamo, and seismic heterogeneities.^{35,36}

Oxygen has been proposed to be a likely light element alloyed with Fe in the liquid OC.⁸ It has been postulated that an oxygen-rich liquid OC adjoining an oxygen-deprived solid IC might account for the global seismic observations of the density jump across the IC–OC boundary.¹¹ The successful syntheses of

Figure 3. Reverse Monte Carlo simulation of Fe₂₈O₁₄ structure (A) Distribution of middle-layer spacings with different neighboring layers. Legends indicate chemical elements in the first and second nearest neighbor layers, e.g., "OFe-FeO" denoting that Fe layers are the first nearest neighbor and that O layers are the second nearest neighbors. The red curve is a Gaussian distribution fitting. (B) Oxygen concentration and XRD deviation (DXRD) of supercell structures from RMC simulation. Each point represents the final structure from one RMC simulation. The color bar indicates the number of atoms in the supercell. (C) The comparison between diffraction patterns by experiments at 220 GPa and the $Fe_{28}O_{14}$ supercell XRD produced by RMC simulations. The incident X-ray wavelength (λ) is 0.2952 Å. Vertical ticks: Fe₂₈O₁₄ (black), B2-FeO (orange), B8-FeO (olive), B2-KCI (magenta), and hcp-Re (blue). The lower panel shows the simulated crystal structure of Fe28014

the solid Fe-rich Fe_nO phases here suggest that the solid IC may also incorporate a significant amount of oxygen. The result is a less drastic oxygen partitioning between the OC and the IC and, consequently, a more complex origin of this density deficit contrast, with other elements also playing a significant role.⁸ In addition, forming solid Fe-rich Fe_nO phases via IC nucleation and growth would affect the associated latent heat release from the IC to power the geodynamo through geological time.³⁷

Oxygen has been present in every stage of Earth's formation and evolution, and chemical interactions in the deep Earth involve oxygen transfer between the core and mantle.^{38–40} It has been suggested that the precipitation of MgO, FeO, and SiO₂ in the OC could have provided gravitational energy to drive early core convection and the Earth's ancient magnetic field before the formation of the Earth's solid

IC.^{41–45} The incorporation of oxygen into solid Fe_nO phase(s) extends the deep oxygen cycle to the IC and impacts this precipitation process. Future studies of multiple element partitioning, including oxygen, between solid and liquid Fe will help clarify the relative abundances of the light elements, the origin of the density deficit in both regions, and many other processes involving the Earth's core evolution.

MATERIALS AND METHODS

Sample preparation and characterization

The starting materials of pure iron (C.A.S. #7439-89-6) mixed with FeO (C.A.S. #1345-25-1) or Fe₂O₃ (C.A.S. #1309-37-1) powders with grain sizes ranging from submicron to a few microns were prepared in an Fe/O ratio ~3:1. Each of the two mixtures was mechanically ground for 90 min in a glove box filled with Ar. Pairs of single- or double-beveled diamond anvils with culet sizes of 50-300 or 60-180-300 μ m were used for the laser-heated DAC experiments. The sample disk of each Fe-O mixture was compressed to ~3- μ m thick and 14-16 μ m across and was selected and loaded into a sample chamber measuring 20-25 μ m across consisting of a pre-indented Re gasket. Each sample platelet was sandwiched between two dried KCl platelets used as the pressure-transmitting medium and thermal insulation layers. The whole DAC sample assembly was sealed in a vacuum using a high-pressure gas loading system after being evacuated for 30 min before each sample chamber was compressed to high pressures. Before reaching a target pressure, XRD patterns were collected at ~30 GPa to confirm the presence of both iron and iron oxide.

Synchrotron XRD experiments

High-P-T XRD measurements were conducted at beamlines 13-IDD of the Advanced Phonon Source, Argonne National Laboratory. A highly monochromatized X-ray source with an energy of 42 keV (0.2952 Å) was used as the incident X-ray beam and focused



The Innovation



Figure 4. Projected DOS of Fe 3d electrons in Fe and Fe_nO compounds at 220 GPa and T_{el} = 3,000 K (A) e-Fe. The dashed line indicates the Fermi level. (B) Fe₃O (space group P6₃/mmc). The blue area shows the partial DOS of Fe ions with oxygen neighbor layers (i.e., Fe1, Fe2, Fe4, and Fe5 shown in the crystal structure), while the magenta area shows the partial DOS of Fe ions without oxygen neighbors (Fe3, Fe6). The purple area shows the partial DOS of O. (C) Fe₂O ($P\overline{3}m1$). (D) B8-FeO. Gold and red spheres represent iron and oxygen ions, respectively. The DOS is normalized by dividing the total number of atoms in the cell.

down to a 2- to 5- μ m spot size (full width at half maximum) at the sample position. Each DAC was compressed to high pressure at room temperature in preparation for the laser heating experiments. The pressure was determined from the lattice parameters of hcp–Fe and B2–KCl, generally with an uncertainty of 2–5 GPa.⁴⁶ At targeted pressures, two infrared laser beams were focused on 20- to 30- μ m spots (full width at half maximum) on both sides of the sample. They were co-axially aligned with the incident X-ray beam using the X-ray-induced luminescence on the sample and/or ruby. We estimated the temperature of the heated samples by fitting measured thermal radiation spectra with the gray-body assumption function. The temperature uncertainty was within 100–200 K based on multiple temperature measurements on both sides.

AGA search

We searched for crystal structures of Fe-rich Fe_nO using the AGA^{28,47,48} which combines *ab initio* calculations and auxiliary interatomic potentials described by the embedded-atom method in an adaptive manner to ensure high efficiency and accuracy. The structure searches were only constrained by the chemical composition, without any assumption on the Bravais lattice type, symmetry, atom basis, or unit cell dimensions. A wide range of different Fe-rich compositions (i.e., 2:1, 3:1, 3:2, 4:1, 4:3, 5:1, 5:2, 5:3, 5:4, 6:1, 6:5, 7:1, 8:1, 9:1) were selected with up to 25 atoms in the unit cell to perform the search. Each AGA run contains a genetic algorithm (GA) loop accelerated by interatomic potential and a density functional theory (DFT) calculation loop to refine the potential. The candidate pool in the GA search contained 64 structures. In each GA generation, 16 new structures were generated from the parent structure pool via the mating procedure. The structure search with a given auxiliary interatomic potential sustained 1000 consecutive GA generations. After one GA loop, 16

structures were randomly selected for the static DFT calculation to refine the potential. This process was repeated 80 times, followed by DFT structure optimization on unique crystal structures.

Ab initio calculations

We conducted ab initio calculations using the projector augmented-wave (PAW) method within DFT, as implemented in the VASP code.^{49–51} The exchange and correlation energies were treated with the generalized gradient approximation parameterized by the Perdew-Burke-Ernzerhof formula.⁵² PAW potentials with valence electronic configurations 3d⁷4s1 and $2s^22p^4$ were used for the Fe and O atoms, respectively. The hard-core PAW potential with 3s²3p⁶3d⁷4s¹ for Fe was also tested. Both Fe PAWs show consistent results (see Note S4 and Figure S17). The Mermin functional was employed.^{53,54} A plane-wave basis set was used with a kinetic energy cutoff of 650 eV. During the AGA search, the Monkhorst-Pack sampling scheme was adopted for Brillouin zone sampling with a k-point grid of $2\pi \times 0.033 \text{ Å}^{-1}$, s and the ionic relaxations stopped when the forces on every atom became smaller than 0.01 eV/Å. The energy convergence criterion is 10⁻⁴ eV. Phonon calculations were performed using density functional perturbation theory implemented in the VASP code and the Phonopy software,^{56,57} with a k-point grid of $2\pi \times 0.033$ Å⁻¹ and supercells of 3 \times 3 \times 2 for Fe₂O (54 atoms), 2 \times 2 \times 2 for Fe₃O (64 atoms), 2 × 2 × 2 for Fe₃O₂ rhombohedral lattice (40 atoms), and 2 × 2 × 2 for Fe₄O₃ rhombohedral lattice (56 atoms).

RMC simulations

The RMC simulations were performed by randomly changing the chemistry (Fe/O ratio) and layer stackings in the supercells of the hexagonal lattice. There is no restriction



Figure 5. Electron localization function (A) ϵ -Fe, (B) Fe₃O (space group *P*6₃/*mmc*), (C and D) Fe₂O (*P*3*m*1), and (D) 88-FeO. Gold and red spheres represent iron and oxygen, respectively.

of the overall Fe/O concentration in the supercell. The only criterion to select preferred structures during the random sampling is based on the XRD deviation $D_{XRD} = \sqrt{(I_{exp}(2\theta) - I_{sim}(2\theta))^2}$, which is the mean-square deviation of the intensities between simulated, I_{sim} , and experimental, I_{exp} , XRDs. The smaller the D_{XRD} , the better the agreement with the experimental data. The acceptance rate during the RMC simulation follows the typical Metropolis acceptance rule as $p = e^{-\beta dD}$, where ΔD is the change of XRD deviation between two MC steps. β is a scaling factor that makes the overall acceptance rate around 20%. The structure is updated if the stacking changes during the RMC simulation. The interlayer distances were set using the most probable spacings shown in Figure 3A. According to Figure 2D, only seven out of 12 stacking sequences appeared in the low-energy structures found with the AGA search. Stackings other than the seven preferred ones in the RMC simulations were discouraged by assigning large ΔD s. The RMC simulations used supercells containing 20 to 60 layers. Each simulation was repeated independently 30 times.

Data availability

6

All original data supporting the findings and structure files of this study have been deposited at https://doi.org/10.5281/zenodo.7155153.

Table 1. Bader charge analysis of the FenO phases

Phase	Fe	0
ε-Fe	0	-
Fe ₃ 0 (P6 ₃ / mmc)	Fe1/Fe2/Fe4/Fe5: +0.45, Fe3/Fe6: +0.01	-0.91
Fe ₂ 0 (<i>P</i> 3 <i>m</i> 1)	+0.47	-0.94
B8-FeO	+0.97	-0.97

The indices of Fe in Fe₃O correspond to those in Figures 4B and 5.

REFERENCES

- Hirose, K., Labrosse, S., and Hernlund, J. (2013). Composition and state of the core. Annu. Rev. Earth Planet Sci. 41, 657–691.
- Birch, F. (1952). Elasticity and constitution of the Earth's interior. J. Geophys. Res. 57, 227–286.
- 3. Poirier, J.-P. (1994). Light elements in the Earth's outer core: a critical review. Phys. Earth Planet. In. **85**, 319–337.
- Hirose, K., Wood, B., and Vočadlo, L. (2021). Light elements in the Earth's core. Nat. Rev. Earth Environ. 2, 645–658.
- Morard, G., Andrault, D., Antonangeli, D., and Bouchet, J. (2014). Properties of iron alloys under the Earth's core conditions. Compt. Rendus Geosci. 346, 130–139.
- Lin, J.-F., Heinz, D.L., Campbell, A.J., et al. (2002). Iron-silicon alloy in Earth's core? Science 295, 313–315.
- Oka, K., Hirose, K., Tagawa, S., et al. (2019). Melting in the Fe-FeO system to 204 GPa: implications for oxygen in Earth's core. Am. Mineral. 104, 1603–1607.
- Badro, J., Brodholt, J.P., Piet, H., et al. (2015). Core formation and core composition from coupled geochemical and geophysical constraints. Proc. Natl. Acad. Sci. USA **112**, 12310–12314.
- Hu, Q., Kim, D.Y., Yang, W., et al. (2016). FeO₂ and FeOOH under deep lower-mantle conditions and Earth's oxygen-hydrogen cycles. Nature 534, 241–244.
- Liu, J., Hu, Q., Bi, W., et al. (2019). Altered chemistry of oxygen and iron under deep Earth conditions. Nat. Commun. 10, 153.
- Alfè, D., Gillan, M.J., and Price, G.D. (2000). Constraints on the composition of the Earth's core from ab initio calculations. Nature 405, 172–175.
- Ozawa, H., Hirose, K., Tateno, S., et al. (2010). Phase transition boundary between B1 and B8 structures of FeO up to 210 GPa. Phys. Earth Planet. In. 179, 157–163.
- Sherman, D.M. (1995). Stability of possible Fe-FeS and Fe-FeO alloy phases at high pressure and the composition of the Earth's core. Earth Planet Sci. Lett. 132, 87–98.
- Alfè, D., Price, G., and Gillan, M.J. (1999). Oxygen in the Earth's core: a first-principles study. Phys. Earth Planet. In. 110, 191–210.
- Weerasinghe, G.L., Pickard, C.J., and Needs, R.J. (2015). Computational searches for iron oxides at high pressures. J. Phys. Condens. Matter 27, 455501.
- Fei, Y., and Mao, H.K. (1994). In situ determination of the NiAs phase of FeO at high pressure and temperature. Science 266, 1678–1680.
- Ozawa, H., Hirose, K., Ohta, K., et al. (2011). Spin crossover, structural change, and metallization in NiAs-type FeO at high pressure. Phys. Rev. B 84, 134417.

- Weck, G., Desgreniers, S., Loubeyre, P., and Mezouar, M. (2009). Single-crystal structural characterization of the metallic phase of oxygen. Phys. Rev. Lett. **102**, 255503.
- Ringwood, A.E., and Hibberson, W. (1990). The system Fe-FeO revisited. Phys. Chem. Miner. 17, 313–319.
- Ozawa, H., Takahashi, F., Hirose, K., et al. (2011). Phase transition of FeO and stratification in Earth's outer core. Science **334**, 792–794.
- Lavina, B., Dera, P., Kim, E., et al. (2011). Discovery of the recoverable high-pressure iron oxide Fe₄O₅. Proc. Natl. Acad. Sci. USA **108**, 17281–17285.
- Lavina, B., and Meng, Y. (2015). Unraveling the complexity of iron oxides at high pressure and temperature: synthesis of Fe₅O₆. Sci. Adv. 1, e1400260.
- Bykova, E., Dubrovinsky, L., Dubrovinskaia, N., et al. (2016). Structural complexity of simple Fe₂O₃ at high pressures and temperatures. Nat. Commun. 7, 10661.
- Liu, J., Wang, C., Lv, C., et al. (2021). Evidence for oxygenation of Fe-Mg oxides at mid-mantle conditions and the rise of deep oxygen. Natl. Sci. Rev. 8, nwaa096.
- Chen, B., Li, Z., Zhang, D., et al. (2014). Hidden carbon in Earth's inner core revealed by shear softening in dense Fe₇C₃. Proc. Natl. Acad. Sci. USA **111**, 17755–17758.
- Prescher, C., Dubrovinsky, L., Liu, X., et al. (2015). High Poisson's ratio of Earth's inner core explained by carbon alloying. Nat. Geosci. 8, 220–223.
- Liu, J., Lin, J., Prakapenka, V.B., et al. (2016). Phase relations of Fe₃C and Fe₇C₃ up to 185 GPa and 5200 K: Implication for the stability of iron carbide in the Earth's core. Geophys. Res. Lett. 43, 12415–12422.
- Wu, S.Q., Ji, M., Wang, C.Z., et al. (2014). An adaptive genetic algorithm for crystal structure prediction. J. Phys. Condens. Matter 26, 035402.
- Zhang, Z., Zhang, D.B., Sun, T., and Wentzcovitch, R. (2019). phq: a Fortran code to compute phonon quasiparticle properties and dispersions. Comput. Phys. Commun. 243, 110–120.
- Sun, Y., Arapan, S., and Rosengren, A. (2022). Ab Initio Melting Temperatures of Bcc and Hcp Iron under the Earth's Inner Core Condition. Preprint at arXiv. https://doi.org/10.48550/arXiv. 2205.02290.
- Zhang, D.-B., Sun, T., and Wentzcovitch, R.M. (2014). Phonon quasiparticles and anharmonic free energy in complex systems. Phys. Rev. Lett. *112*, 058501.
- Becke, A.D., and Edgecombe, K.E. (1990). A simple measure of electron localization in atomic and molecular systems. J. Chem. Phys. 92, 5397–5403.
- Silvi, B., and Savin, A. (1994). Classification of chemical bonds based on topological analysis of electron localization functions. Nature 371, 683–686.
- Henkelman, G., Arnaldsson, A., and Jónsson, H. (2006). A fast and robust algorithm for Bader decomposition of charge density. Comput. Mater. Sci. 36, 354–360.
- Hsieh, W.-P., Goncharov, A.F., Labrosse, S., et al. (2020). Low thermal conductivity of iron-silicon alloys at Earth's core conditions with implications for the geodynamo. Nat. Commun. 11, 3332.
- Zhang, Y., Hou, M., Liu, G., et al. (2020). Reconciliation of experiments and theory on transport properties of iron and the geodynamo. Phys. Rev. Lett. **125**, 078501.
- O'Rourke, J.G., Korenaga, J., and Stevenson, D.J. (2017). Thermal evoution of Earth with magnesium precipitation in the core. Earth Planet Sci. Lett. 263–272.
- Knittle, E., and Jeanloz, R. (1991). Earth's core-mantle boundary: results of experiments at high pressures and temperatures. Science 251, 1438–1443.
- Takafuji, N., Hirose, K., Mitome, M., and Bando, Y. (2005). Solubilities of O and Si in liquid iron in equilibrium with (Mg, Fe)SiO₃ perovskite and the light elements in the core. Geophys. Res. Lett. **32**, L06313.
- Tsuno, K., Frost, D.J., and Rubie, D.C. (2013). Simultaneous partitioning of silicon and oxygen into the Earth's core during early Earth differentiation. Geophys. Res. Lett. 40, 66–71.
- O'Rourke, J.G., and Stevenson, D.J. (2016). Powering Earth's dynamo with magnesium precipitation from the core. Nature 529, 387–389.
- Badro, J., Siebert, J., and Nimmo, F. (2016). An early geodynamo driven by exsolution of mantle components from Earth's core. Nature 536, 326–328.
- Hirose, K., Morard, G., Sinmyo, R., et al. (2017). Crystallization of silicon dioxide and compositional evolution of the Earth's core. Nature 543, 99–102.
- Davies, C.J., Pozzo, M., Gubbins, D., and Alfè, D. (2020). Transfer of oxygen to Earth's core from a long-lived magma ocean. Earth Planet Sci. Lett. 538, 116208.
- Zhang, Z., Csányi, G., Alfè, D., et al. (2022). Free energies of Fe-O-Si ternary liquids at high temperatures and pressures: implications for the evolution of the Earth's core composition. Geophys. Res. Lett. 49, e2021GL096749.
- Dewaele, A., Belonoshko, A.B., Garbarino, G., et al. (2012). High-pressure-high-temperature equation of state of KCI and KBr. Phys. Rev. B 85, 214105.
- Deaven, D.M., and Ho, K.M. (1995). Molecular geometry optimization with a genetic algorithm. Phys. Rev. Lett. 75, 288–291.
- Zhao, X., Nguyen, M.C., Zhang, W., et al. (2014). Exploring the structural complexity of intermetallic compounds by an adaptive genetic algorithm. Phys. Rev. Lett. 112, 045502.
- Blöchl, P.E. (1994). Projector augmented-wave method. Phys. Rev. B Condens. Matter 50, 17953–17979.

- Kresse, G., and Furthmüller, J. (1996). Efficiency of ab-initio total energy calculations for metals and semiconductors using a plane-wave basis set. Comput. Mater. Sci. 6, 15–50.
- Kresse, G., and Furthmüller, J. (1996). Efficient iterative schemes for ab initio total-energy calculations using a plane-wave basis set. Phys. Rev. B Condens. Matter 54, 11169–11186.
 Perdew, J.P., Burke, K., and Ernzerhof, M. (1996). Generalized gradient approximation made
- simple. Phys. Rev. Lett. **77**, 3865–3868. 53. Mermin, N.D. (1965). Thermal properties of the inhomogeneous electron gas. Phys. Rev.
- Merrian, N.D. (1903). Therman properties of the inhomogeneous electron gas. Phys. Rev. 137, A1441–A1443.
 Merrian D.M. Merting, H., and Alles, D.D. (1002). Economical frequencies for a section.
- Wentzcovitch, R.M., Martins, J.L., and Allen, P.B. (1992). Energy versus free-energy conservation in first-principles molecular dynamics. Phys. Rev. B Condens. Matter 45, 11372–11374.
- Monkhorst, H.J., and Pack, J.D. (1976). Special points for Brillouin-zone integrations. Phys. Rev. B 13, 5188–5192.
- Baroni, S., de Gironcoli, S., Dal Corso, A., and Giannozzi, P. (2001). Phonons and related crystal properties from density-functional perturbation theory. Rev. Mod. Phys. 73, 515–562.
- 57. Togo, A., and Tanaka, I. (2015). First principles phonon calculations in materials science. Scr. Mater. *108*, 1–5.
- Wentzcovitch, R.M., Yu, Y.G., and Wu, Z. (2010). Thermodynamic properties and phase relations in mantle minerals investigated by first principles quasiharmonic theory. Rev. Mineral. Geochem. 71, 59–98.
- Qin, T., Zhang, Q., Wentzcovitch, R.M., and Umemoto, K. (2019). qha: a Python package for quasiharmonic free energy calculation for multi-configuration systems. Comput. Phys. Commun. 237, 199–207.
- 60. Palasyuk, O., Onyszczak, M., Kim, T.H., et al. (2021). Structural and magnetic properties of hard magnetic system Ce(Co_{1-x}Fe_x)_{4.4}Cu_{0.6} ($0 \le x \le 0.19$). J. Alloys Compd. **883**, 160866.
- Alfè, D., Gillan, M.J., and Price, G.D. (2007). Temperature and composition of the Earth's core. Contemp. Phys. 48, 63–80.

ACKNOWLEDGMENTS

This work is supported by NSFC grants nos. 42072052 and U1930401 and by National Science Foundation awards EAR-1918126 (R.M.W. and Y.S.), EAR-1918134 (K.H. and C.W), and EAR-1901808 and EAR-1916941 (J. Lin). R.M.W. also acknowledges partial support from the Department of Energy, Theoretical Chemistry Program, through grant DE-SC0019759. Computational resources were provided by the Extreme Science and Engineering Discovery Environment (XSEDE) funded by the National Science Foundation through award ACI-1548562. This research also used resources of the Advanced Photon Source, a US Department of Energy Office of Science User Facility operated by Argonne National Laboratory under contract no. DE-AC02-06CH11357. The GeoSoilEnviroCARS at the Advanced Photon Source is supported by the National Science Foundation - Earth Sciences award EAR-1634415 and the Department of Energy-GeoSciences award DE-FG02-94ER14466. Some experiments are supported by the Synergic Extreme Condition User Facility (SECUF) and beamline BL15U1 at the Shanghai Synchrotron Radiation Facility (SSRF).

AUTHOR CONTRIBUTIONS

J. Lin, J. Liu, R.M.W., and Y.S. conceived the project. Y.S., F.Z., C.W, and K.H. performed *ab initio* calculations and crystal structure search. J. Liu, J. Lin, C.L., S.F., and V.B.P. carried out synchrotron experiments and data analyses. J. Liu, Y.S., and R.M.W. wrote the manuscript. All authors discussed the results and contributed to paper writing.

DECLARATION OF INTERESTS

The authors declare no competing interests.

SUPPLEMENTAL INFORMATION

Supplemental information can be found online at https://doi.org/10.1016/j.xinn.2022.100354.

LEAD CONTACT WEBSITE

https://www.jsg.utexas.edu/lin http://mineralscloud.org/

SUPPORTING CITATIONS

The following references appear in the Supplemental Information: ^{58,59,60,61}.