## **Evaluating Climate Projection for Drought and Extreme Surface Temperatures over South-Central US**

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#### 2009 Report on global climate change impacts in the United States (Karl et al)

Projected number of days when

Projected End-of-Century under

Lower Emissions Scenario<sup>91</sup> (2080-2099 Average)

T\_>90F by 2080-2099

IPCC AR4 models projected 15-30% decrease of rainfall and nearly double the number of days when T>90F.

Projected Spring rainfall change by 2080s-2090S compared to that of 1971-2000



#### However,

- Large inter-model discrepancy in projected future rainfall changes
- > Which projections should we believe?



# How can we determine creditability of the CMIP5 climate projection?

- Does the multi-models ensemble projection necessarily outperform individual model projection over Texas and SC US?
  - Gleckler et al. (2008), Pierce et al. (2009): An ensemble mean, especially a multi-model ensemble mean projection, can outperform the best quality model because the former allows cancellation of offsetting errors in the individual global models.
  - > What should we do if majority of the models have similar biases?

### Criteria for our process-based model evaluation Metrics:

Response to increase of the > Relevant to climate global sea surface temperature projection Surface water budget and drought indices > Capture processes that control droughts Surface meteorological conditions over Texas Large-scale circulation > Can be compared to long-term observations **Connection with ENSO** 

## **IPCC AR5 Models and Datasets Used for Evaluation:**

#### Datasets:

- CPC US-Mexicao daily rainfall (Higgins et al. 1996), 1°,
- > GHCN daily Tmax, Tmin (Vose et al. 1992), 2.5°
- > NLADAS (Rodell et al. 2004), ET, 1/8°, 1980-2007.
- > ERSSTv3b SST (Smith et al. 2008), 2.0°, 1854-2005
- NCEP reanalysis (Kalney et al 1996; Kistler et al. 2001), 2.5°, 1948-present

All the datasets and models are re-mapped to 2.5° spatial resolution

#### **Periods:**

- > 1950-2005; meteorological data
- > 1980-2005: surface energy/water balance.



#### Table 1. Description of CMIP5 models used in this study

Model (Fig marker)	Institute (Country)	Available Ensembles	Components (Resolutions)	Calendar	Reference
CCSM4 (A)	National Center for Atmospheric Research (USA)	6	F09_g16 (0.9×1.25_gx1v6)	No leap	Gent et al., 2011
GFDL- ESM2M (B)	NOAA/Geophysical Fluid Dynamics Laboratory (USA)	1	Atm: AM2 (AM2p14, M45L24) Ocn: MOM4.1 $(1.0^{\circ}$ lat $\times 1.0^{\circ}$ lon, enhanced tropical resolution: 1/3 on the equator)	No leap	John Dunne et al., 2012
GFDL- ESM2G (C)	NOAA/Geophysical Fluid Dynamics Laboratory (USA)	1	Atm: AM2 (AM2p14, M45L24) Ocn: MOM4.1 ( $1.0^{\circ}$ lat $\times 1.0^{\circ}$ lon, enhanced tropical resolution: $1/3$ on the equator)	No leap	John Dunne et al., 2012
GISS-E2-R (D)	NASA/Goddard Institute for Space Studies (USA)	5	Atm: GISS-E2 (2.0° lat ×2.5° lon) Ocn: R	No leap	Schmidt et al., 2006
HadGEM2- CC (E)	Met Office Hadley Centre (UK)	3	Atm: HadGAM2 (N96L60) Ocn: HadGOM2 (Lat: 1.0-0.3 Lon: 1.0 L40)	360 d/y	Collins et al., 2011; Martin et al., 2011
MPI-ESM- LR (F)	Max Planck Institute for Meteorology (Germany)	3	Atm: ECHAM6 (T63L47) Ocn: MPIOM (GR15L40)	Gregorian	Raddatz et al., 2007; Marsland et al., 2003
IPSL- CM5A-LR (G)	Institut Pierre Simon Laplace (France)	5	Atm: LMDZ4 (96×95×39, 1.875° lat ×3.75° lon) Ocn: ORCA2 (2×2L31, 2.0° lat ×2.0° lon)	No leap	Marti et al., 2010
MIROC5 (H)	AORI, NIES & JAMSTEC (Japan)	4	Atm: AGCM6 (T85L40) Ocn: COCO (COCO4.5)	No leap	Watanabe et al., 2010
MRI- CGCM3	Meteorological Research Institute	3	Atm: GSMUV (TL159L48)	Gregorian	Yukimoto et al., 2011

### **Evaluate seasonal cycles of climatic surface conditions:**

- Cold bias in daily maximum surface temperature (Tmax)
- Wet biases in Precipitation (P), Evapotranspiration (ET), esp. during spring & summer
- > Large discrepancies in seasonal rainfall



Black line: observations, Bold Red line: multi-model ensemble mean

#### **Probability distributions of** Tmax, Tmin, P and drought indices (SPI6 and SPI9)

- Tmax: underestimate warmer Tmax  $\triangleright$ and overestimate cooler Tmax
- Tmin: underestimate cooler Tmin, overestimate warmer Tmin (consistent with wet bias)
- P: underestimate non-rain and heavy  $\triangleright$ rainrate, overestimate light rainrate
- SPI: reasonably realistic, but  $\geq$ underestimate intensity of extreme drought.







### *Number of days/yr when T<sub>max</sub>>90F & 100F:*

- > Reverse the E-W gradient of extreme Tmax over Texas,
- Most of models overestimate occurrence of extreme Tmax over the southeastern Great Plains,
- > Large inter-model discrepancies







### **Evaluation of Large-scale atmospheric circulation:**

- Most of the models underestimate the 500hPa ridge over central US in summer and strength of jet in spring (except for CCSM4).
- Probably responsible for wet and cold biases in spring and summer.



Figure 6: Comparison of the modeled Z500hPa pattern by each CMIP5 models with that of NCEP-CDAS1.

\*Circles highlight better models

- I/2 models underestimate lower tropospheric westerly winds (U850) in spring and summer.
- Underestimate lower tropospheric southerly winds (V850) in spring





#### Correlation between SC US rainfall anomalies and Nino3 and Nino4 indices:

#### About a half of the models

- > underestimate correlation with ENSO in winter
- > overestimate ENSO connection in spring, summer and fall
- > Because of errors in ENSO teleconnection pattern (not shown)

Figure 9: Correlations between Niño4, Niñ3 and SC US rainfall. "Star" indicates significant correlation coefficient at 95% confidence level using student t-test.



Leading REOF of global SST variance during 1900-2005:

- Observation shows the global increase of sea surface temperature (SST) as the leading mode for SST variance (Schubert et al. 2008).
- Few models realistically capture this global increase of SST mode (CCSM4 and MPI)





# Modeled response of summer rainfall over SC US to the increasing global SST mode:

- Most of the models underestimate the change of summer rainfall over SC US associated with global increase of SST over th period of 1900-2005.
- Only CCSM4 captures the observed relationship between the increase of global SST mode and increase of summer rainfall over SC US.





#### Ranking the models using our process-based metrics:

CCSM4 overall ranks the best, especially in SC US rainfall response to increase of global SST.

Response to increase of the global sea surface temperature

#### Surface conditions

Surface water budget and drought indices

Large-scale circulation

**Connection with ENSO** 

	Table 2: Ranking of model performance for SC US regional climate change											
	Variables Models											
		CCSN	A4 GFDL-	GFDL-	GISS-	HadGE	MPI	IPSL	MIROC	MR		
			ESM2G	ESM2M	E2-R	M2			5	Ι		
Tier-1: Forced variability or change												
	Correlation with global SST warming:											
	a <sub>GW</sub>	1	3	1	3	2	3	3	3	3		
	GW <sub>SST</sub>	2	1	3	1	3	2	2	3	2		
7	Subtotal	1.5	2	2	2	2.5	2.5	2.5	3	2.5		
	Seasonal cycle:											
	Tmax	1	2	2	2	1	2	3	1	2		
1	Tmin	2	1	1	1	3	1	3	2	1		
-	q	1	1	2	1	3	1	3	1	1		
	Subtotal	1.3	1.3	1.7	1.3	2.3	1.3	3	1.3	1.3		
	PD <sub>Tmax</sub>	3	3	3	3	3	3	3	3	2		
	PD <sub>RR</sub>	2	2	2	2	2	2	2	2	1		
	Р	1	3	3	2	3	1	2	2	3		
7	ET	3	2	2	3	2	2	2	2	2		
	SPI6	2	2	2	2	2	2	2	2	2		
	SPI9	2	2	2	2	2	2	2	2	2		
	Subtotal	2.2	2.3	2.3	2.3	2.3	2	2.2	2.2	2		
	Z500	2	3	3	3	2	2	3	2	3		
7	U850	1	2	2	2	2	1	1	2	2		
	V850	2	2	2	2	1	2	2	2	2		
	Subtotal	1.7	2.3	2.3	23	1.7	1.7	2	2	2.3		
	Tier-2: natural variability											
	r <sub>p,Niño3</sub>	3	2	2	1	3	3	3	2	3		
	SZ500,	2	2	2	3	3	3	3	3	3		
	NIN03			2		-		2				
	r <sub>p,Niño4</sub>	3	2	2	1	3	3	3	2	3		
_	SZ500,	2	2	2	3	3	3	3	2	3		
1	Niño4	2.5	2	2	2	2	2	2	2.2	2		
	Subtotal	2.5	2	2	2	3	3	3	2.3	3		

#### Projected change of Tmax during 2073-2099 relative to 1979-2005:

- Models consistently project a disproportional increase of occurrence of high Tmax (>90F -108F) by
  - 25-50% under low emission (but unlikely RCP4.5) scenario (CO<sub>2</sub> reaches 650 ppm by 2100)
  - 50-100% under high emission (business as usual, RCP8.5) scenario (CO<sub>2</sub> reaches 1350ppm by 2100)
- Recall that these models tend to underestimate Tmax.



#### Projected change of Tmin in 2073-2099 relative to 1979-2005.

➤ Models consistently project a strong increase of occurrence of Tmin≥80F several folds under the high emission (RCP8.5) scenario.



Daily Minimum Surface Temperature Probability Distribution Function (K)

#### Projected change of rainrate in 2073-2099 relative to 1979-2005.

Increase of nonrainy days and low rainrate and decrease of medium rainrate.



Rainfall Probability Distribution Function (mm/day)

# Projected change of surface net water flux in 2073-2099 relative to 1979-2005:

Under the high emission (business as usual, RCP8.5) scenario:

- Both multi-models and best performing model project net drying, by ~20% of P-ET in spring and summer, despite differences in details.
- Increase of rainfall (P) and ET during winter and spring, decrease of rainfall and ET in summer.
- Net drying in spring is dominated by increase of ET, whereas drying in summer is dominated by decrease <u>A(P-ET)</u> of P.
- > Outliners in projections tends to be the worst performing models.



# **Conclusions:**

The 9 climate models that participated in the IPCC AR5 we evaluated

- share common wet and cold biases, due to underestimate mid-tropospheric ridge in summer, the upper-level jet strength and westerly low-level winds in spring. Most of the models cannot adequately capture the changes of SC US rainfall with ENSO and the increase of global SST.
- consistently project ~20% decrease of net P-ET (dry) in spring-summer by 2073-2099 relative to 1979-2005, under the "business as usual" emission scenario (RCP8.5), despite differences in details.