

The role of the Mexican Plateau in shaping rainfall over Texas

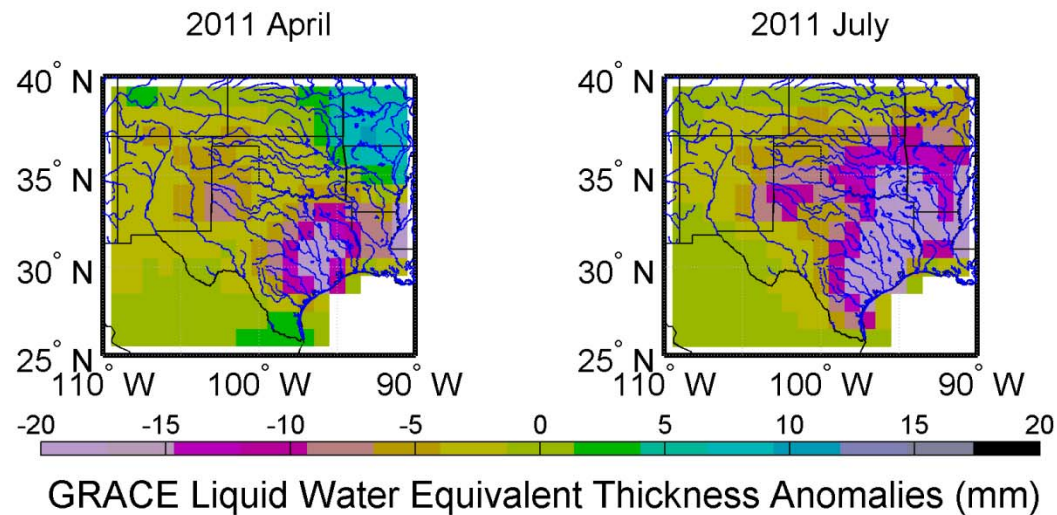
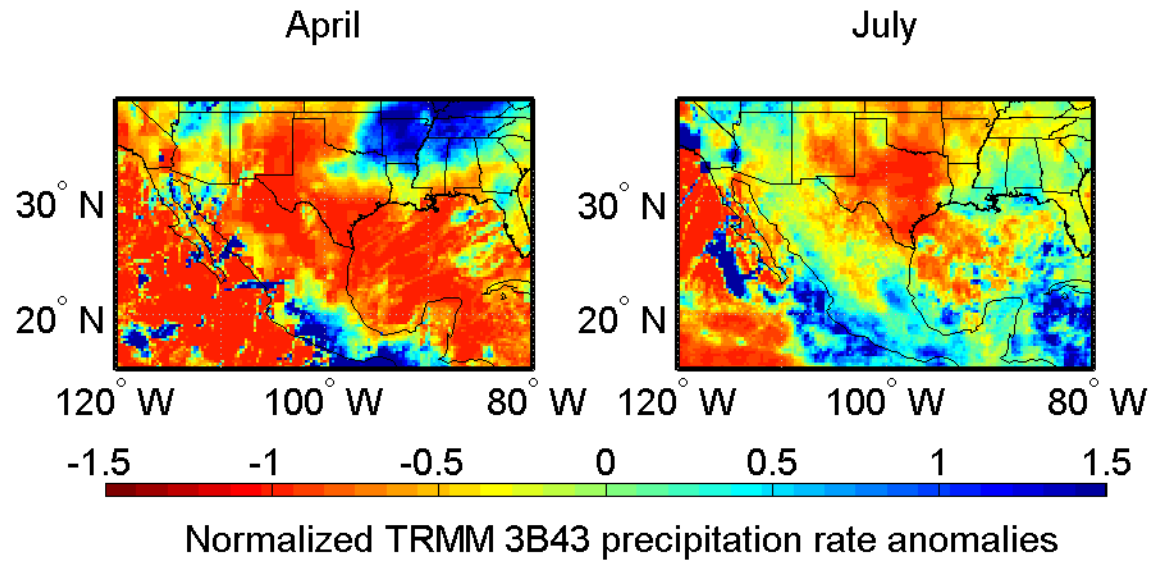
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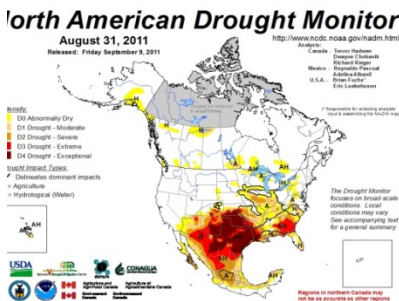
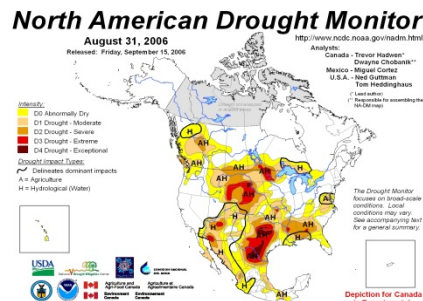
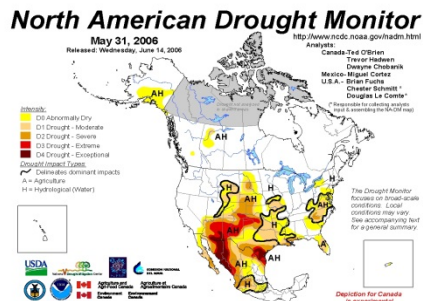
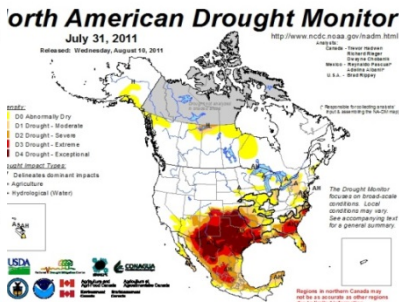
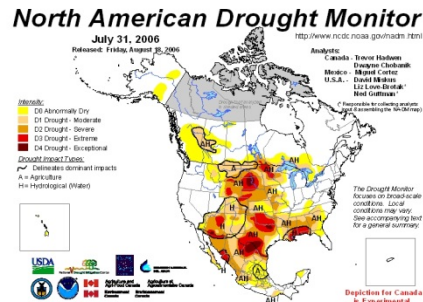
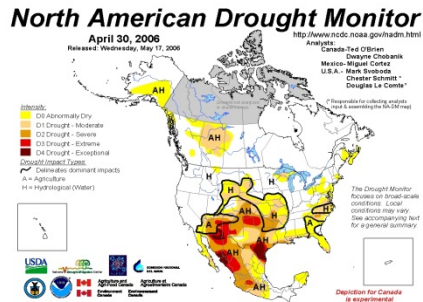
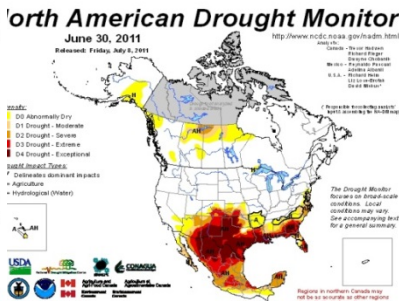
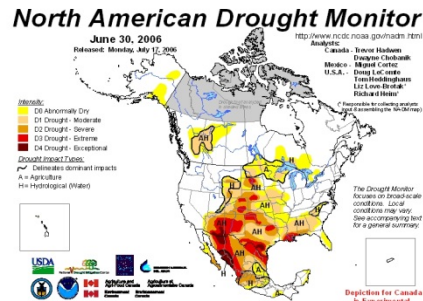
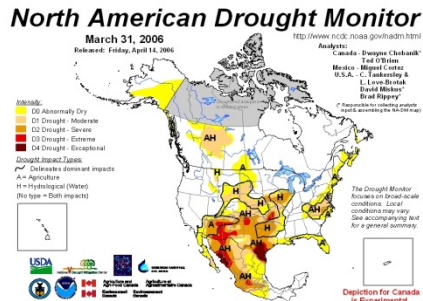
Rainfall and soil moisture deficit



North American Drought Monitor

2006 spring and summer

2011 spring and summer



Background

SST Anomalies Force? → Texas drought

- A drought generally results from a synthesis of numerous factors (e.g. Hoerling et al. 2013)
- The drought continuation into the summer of 2011 was not significantly SST-forced (Seager et al. 2013)

Mexican Plateau drought

Exacerbate?

- Warm dry air advection in spring and summer (Lanicci et al. 1987, Myoung and Nielsen-Gammon 2010)
- Air descends over Texas in summer (Barlow et al. 1998)

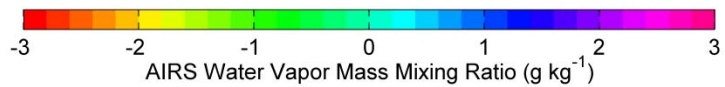
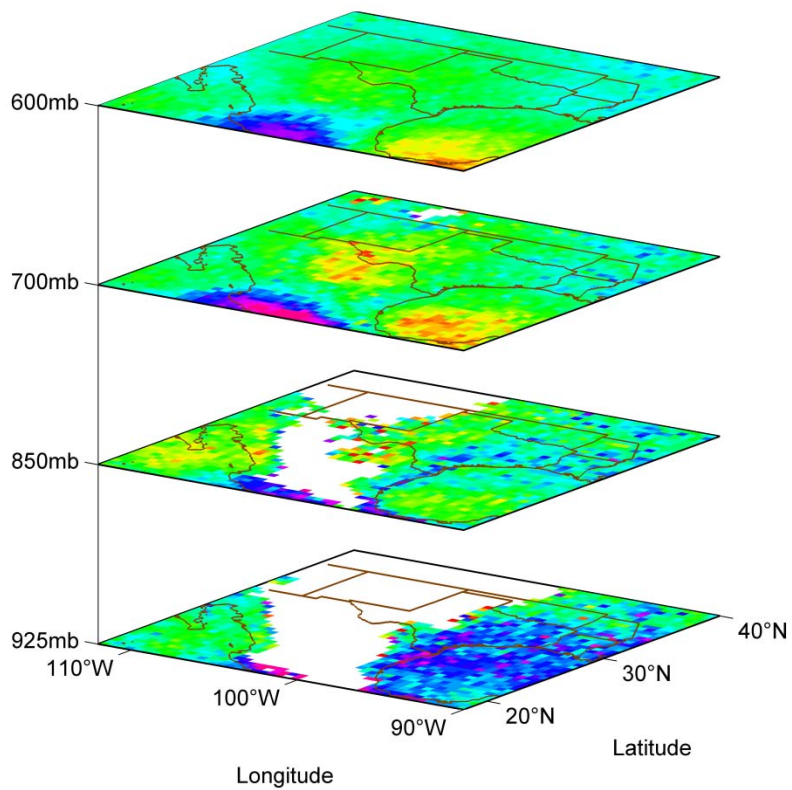
- Global warming (e.g. Hoerling et al. 2013)
- Soil moisture-precipitation feedback (e.g. Su et al. 2013)
- Aerosols?

Other factors

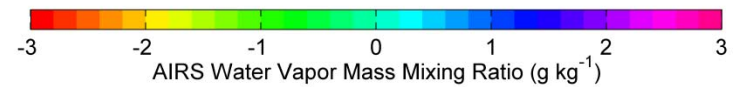
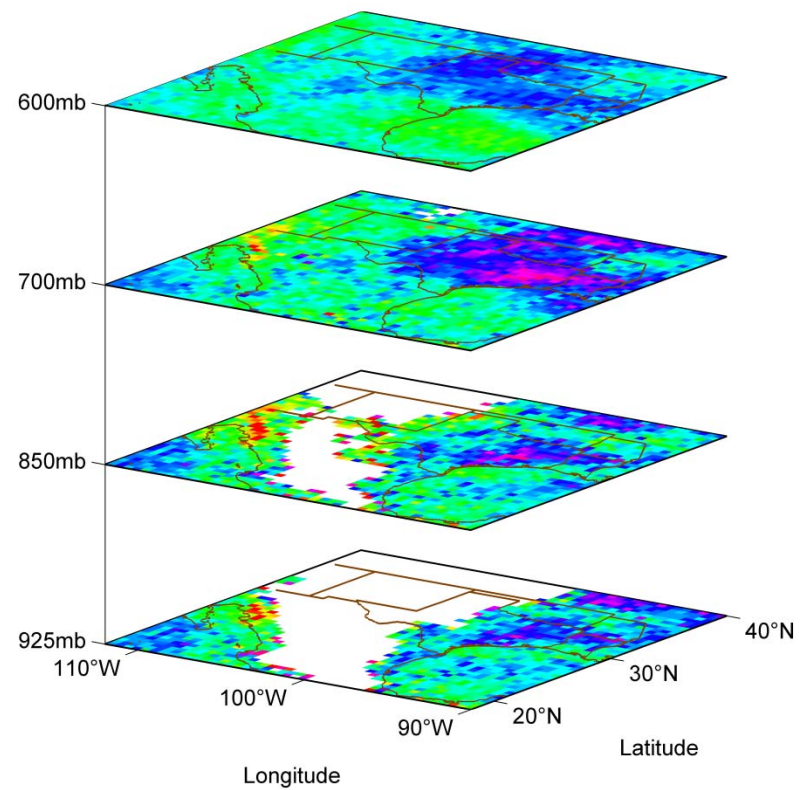


Differences in low level water vapor content

2011 April – 2010 April

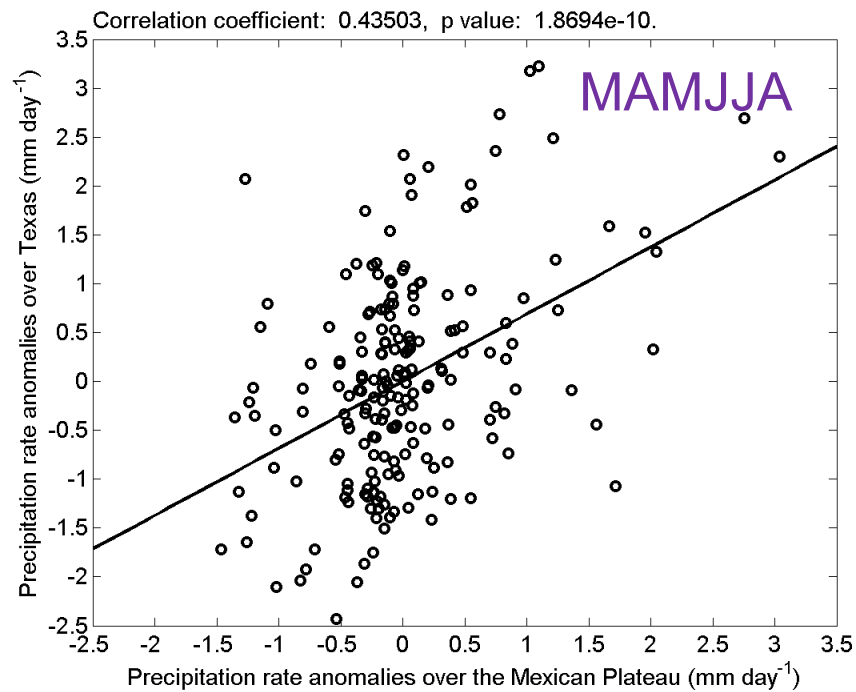


2011 July – 2010 July



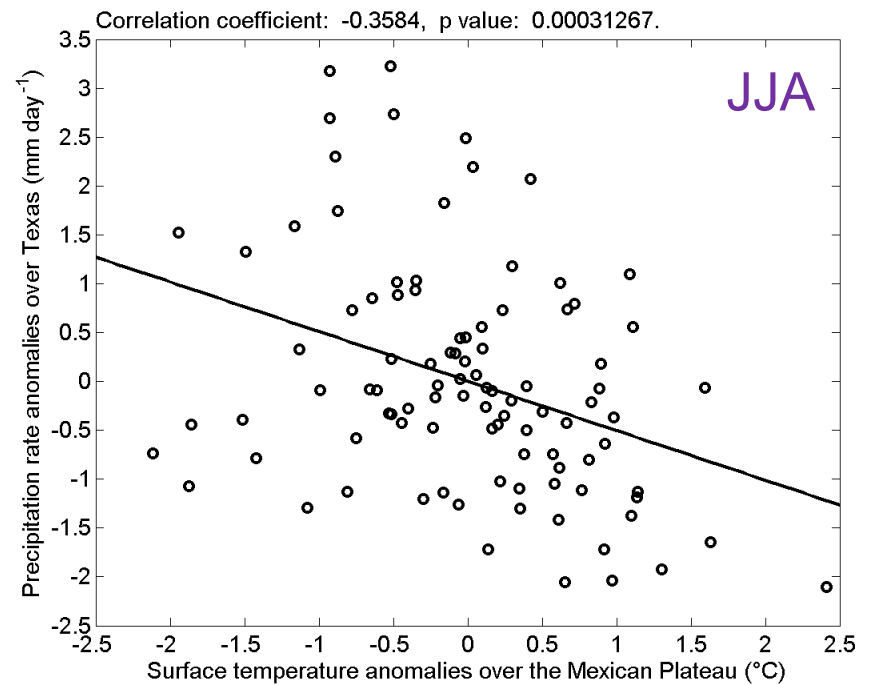
Statistical relationships between the surface temperature/precipitation rate of the Mexican Plateau and the precipitation rate over Texas

Precipitation .vs. Precipitation



$$R = 0.44$$

Precipitation .vs. Temperature

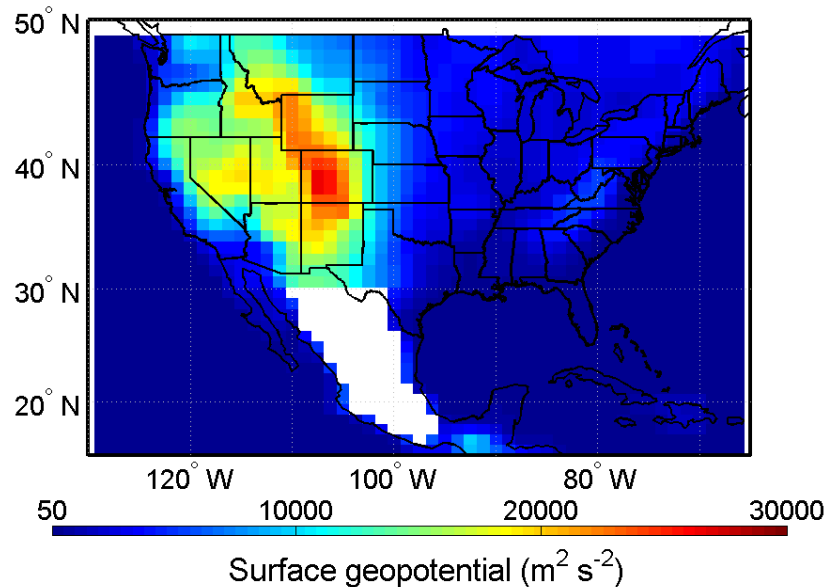


$$R = -0.36$$

Hypothesis and experiment design

Hypothesis:

During the summer, a warmer Mexican Plateau tends to bend the low-level jet towards the highlands and thus an anti-cyclonic flow anomaly forms over the southern US, which tend to diverge the air and reduce rainfall.

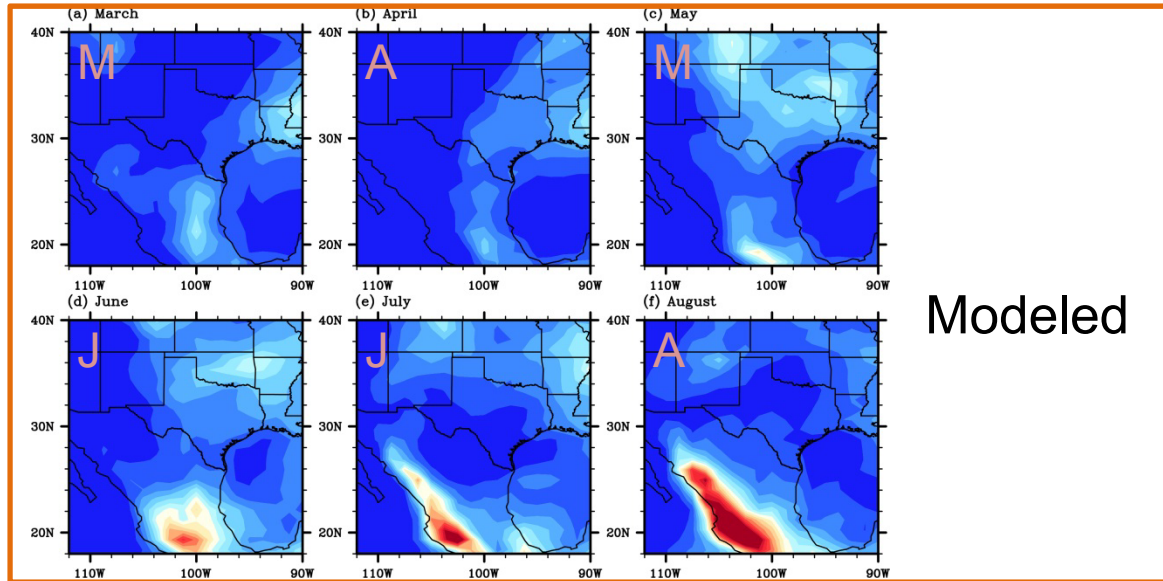


- Prescribed climatological SST
- Integrated 12 years
- At a horizontal resolution of 1° × 1°
- Discarded 1st year for spin up

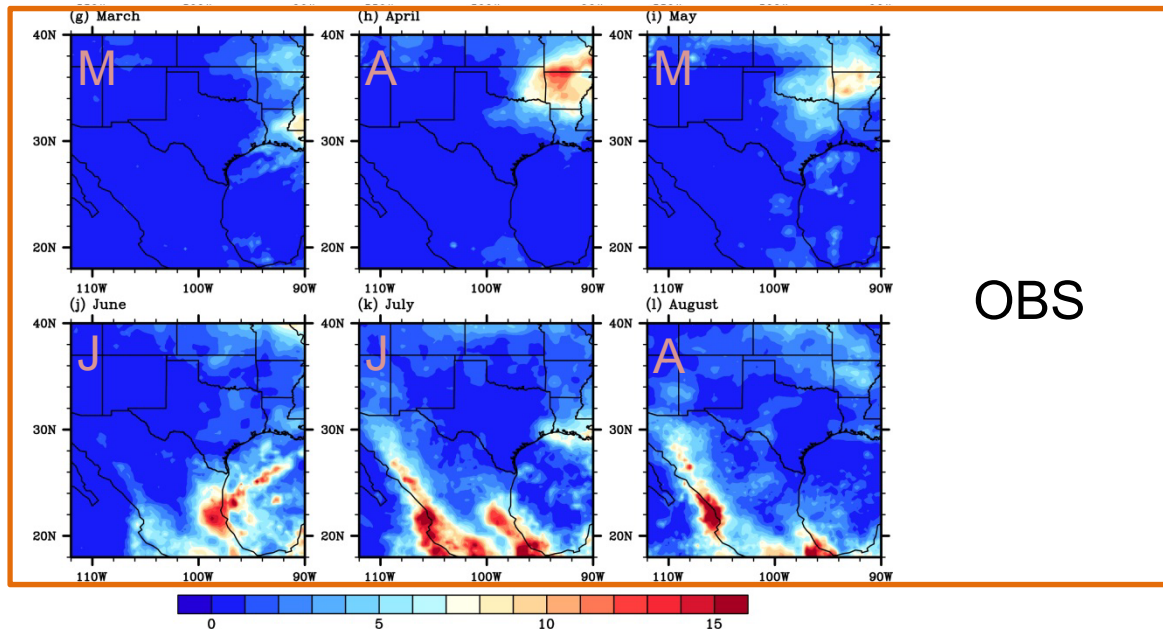
	Topography	Albedo	Soil water
Control run (CTRL)	Default	Default	Default
Experiment 1 (E1 DRY)	Default	Default	0.2 mm for each layer
Experiment 2 (E2 WARM)	Default	0	Default
Experiment 3 (E3 DRY & WARM)	Default	0	0.2 mm for each layer

CESM .vs. observations I with prescribed observed SST

Precipitation rate (mm/day)



*Precipitation rate
(mm/day) in 2011*

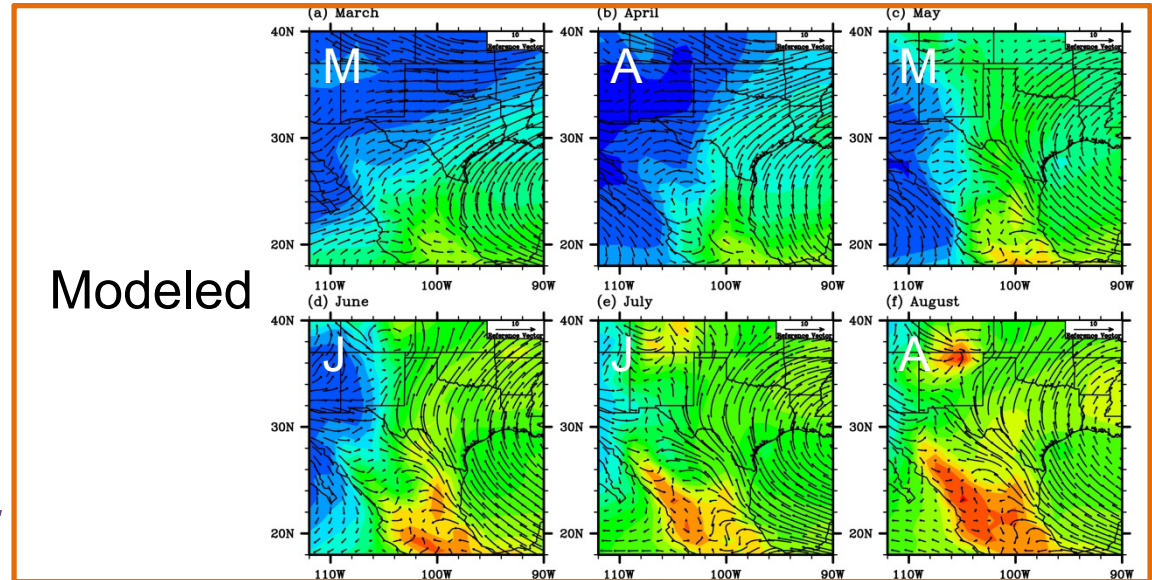


CESM .vs. observations II with prescribed observed SST

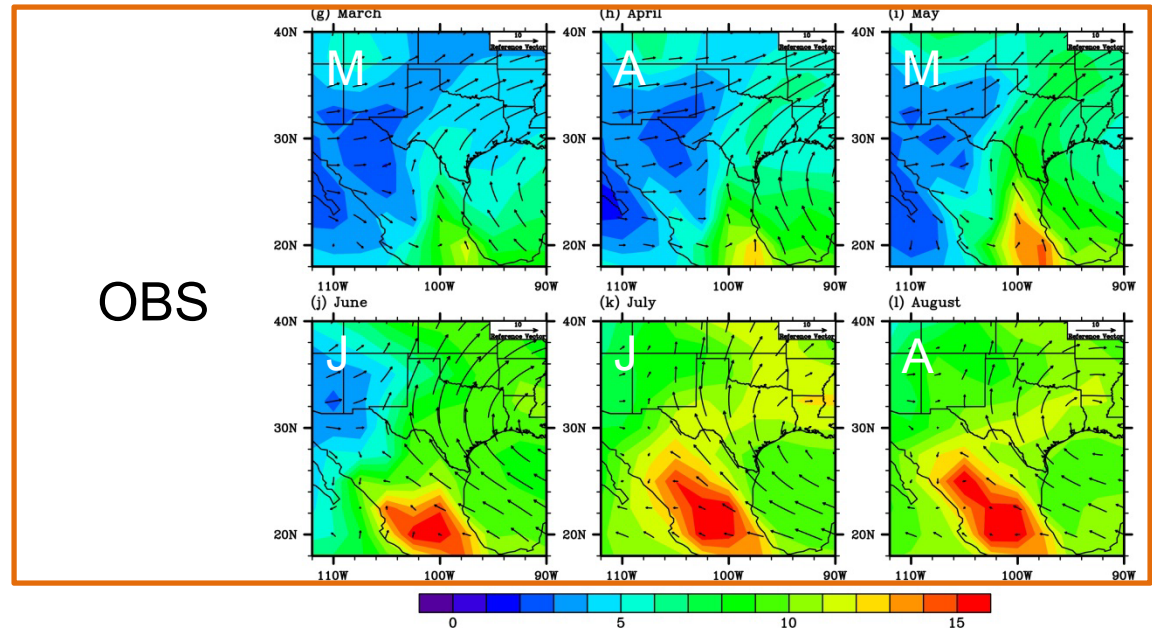
850 mb specific humidity (g/kg) and horizontal wind (m/s)

*Specific humidity (g/kg)
and wind speed (m/s) at
850 hPa in 2011*

Modeled

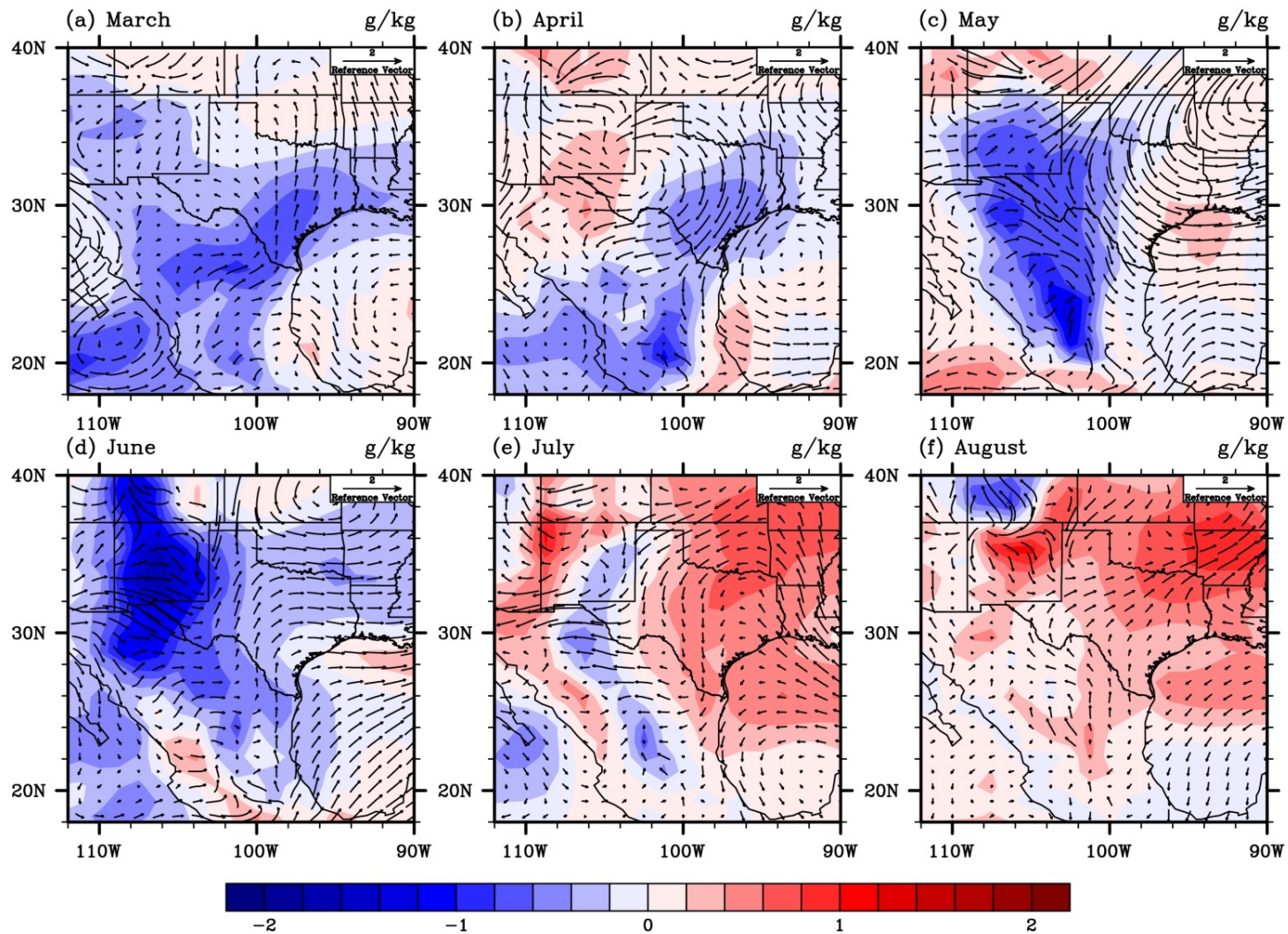


OBS



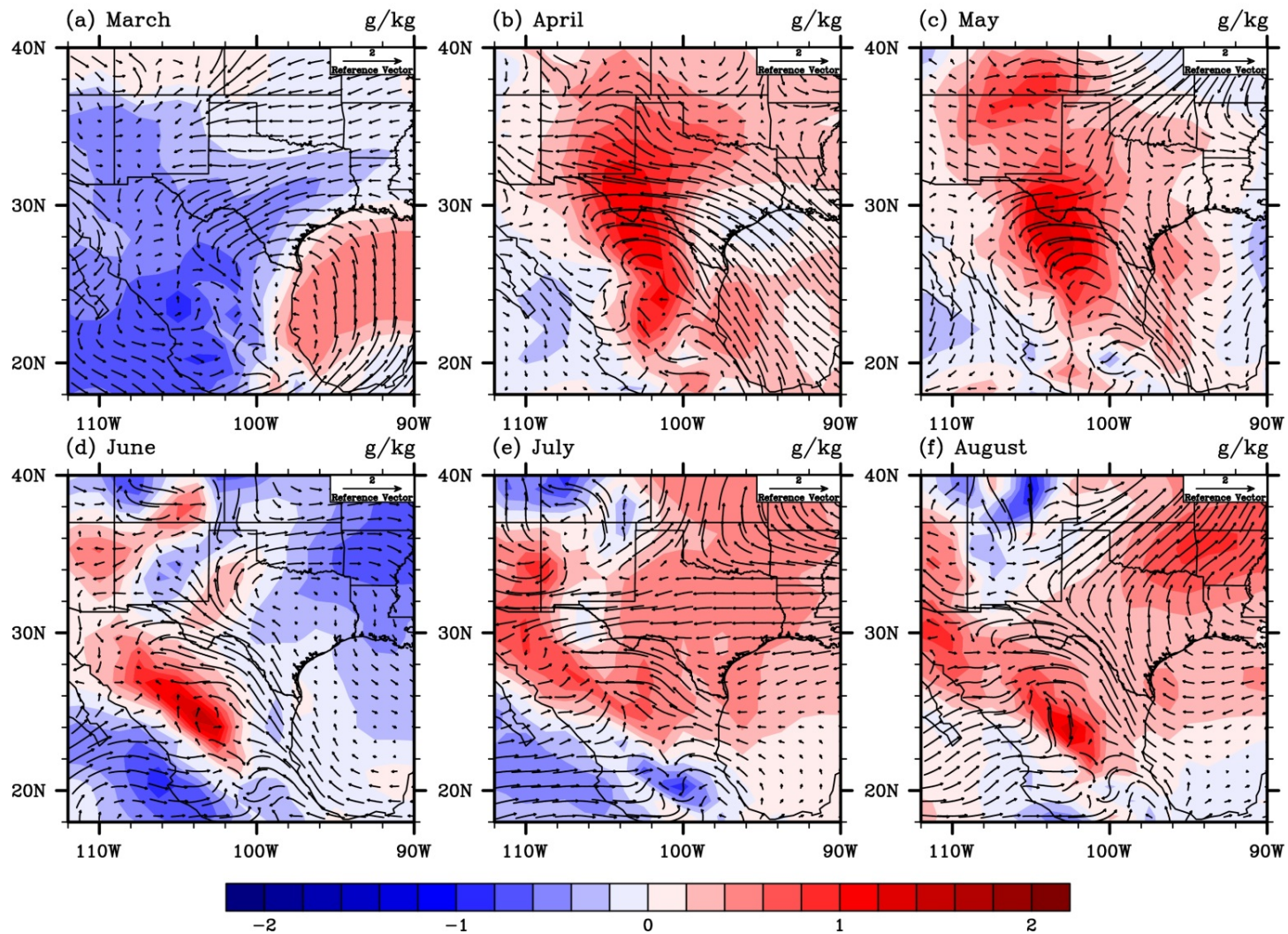
850 hPa specific humidity decreases over the Mexican highlands and the downstream regions

E1(DRY) minus CTRL



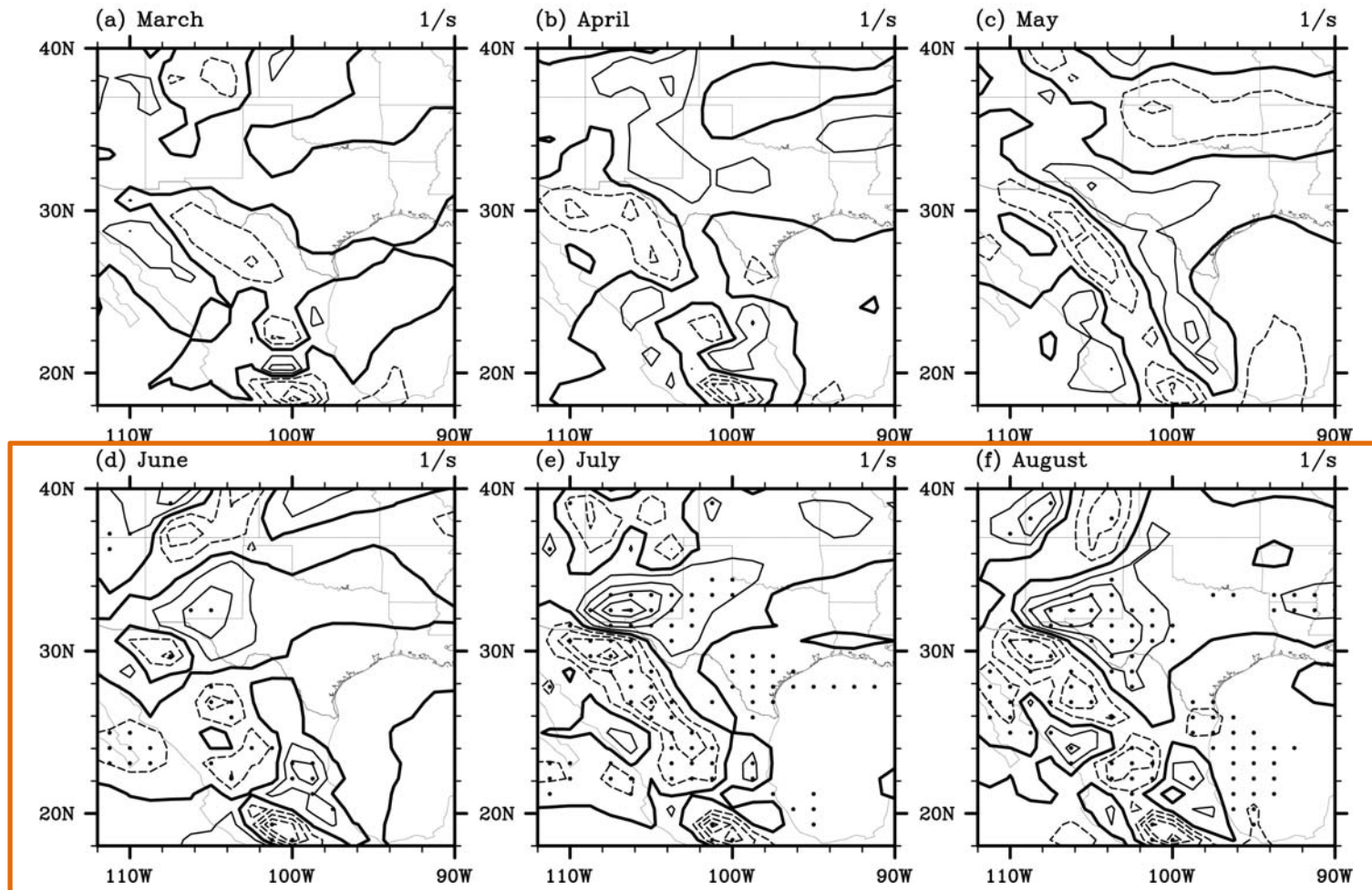
A warm low anomaly over the Mexican Plateau and a summer anti-cyclonic flow anomaly over Texas

E2(WARM) minus CTRL



850 hPa air divergence

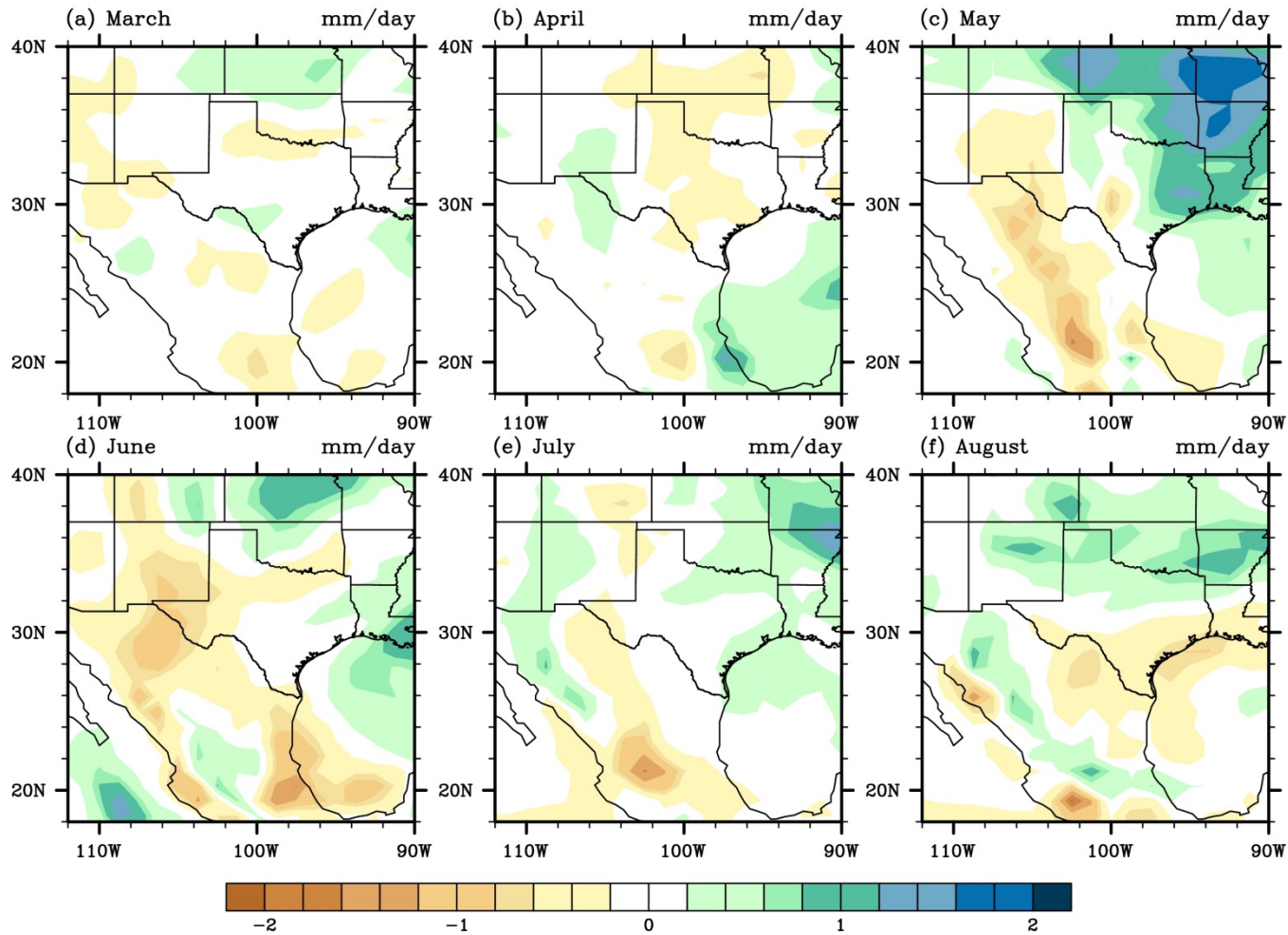
E2(WARM)-CTRL-850hPa Air Divergence



The solid contours represent the air divergence tendency whereas the dashed contours represent the air convergence tendency. The dots represent the areas that pass a student's t-test at the 95% confidence level.

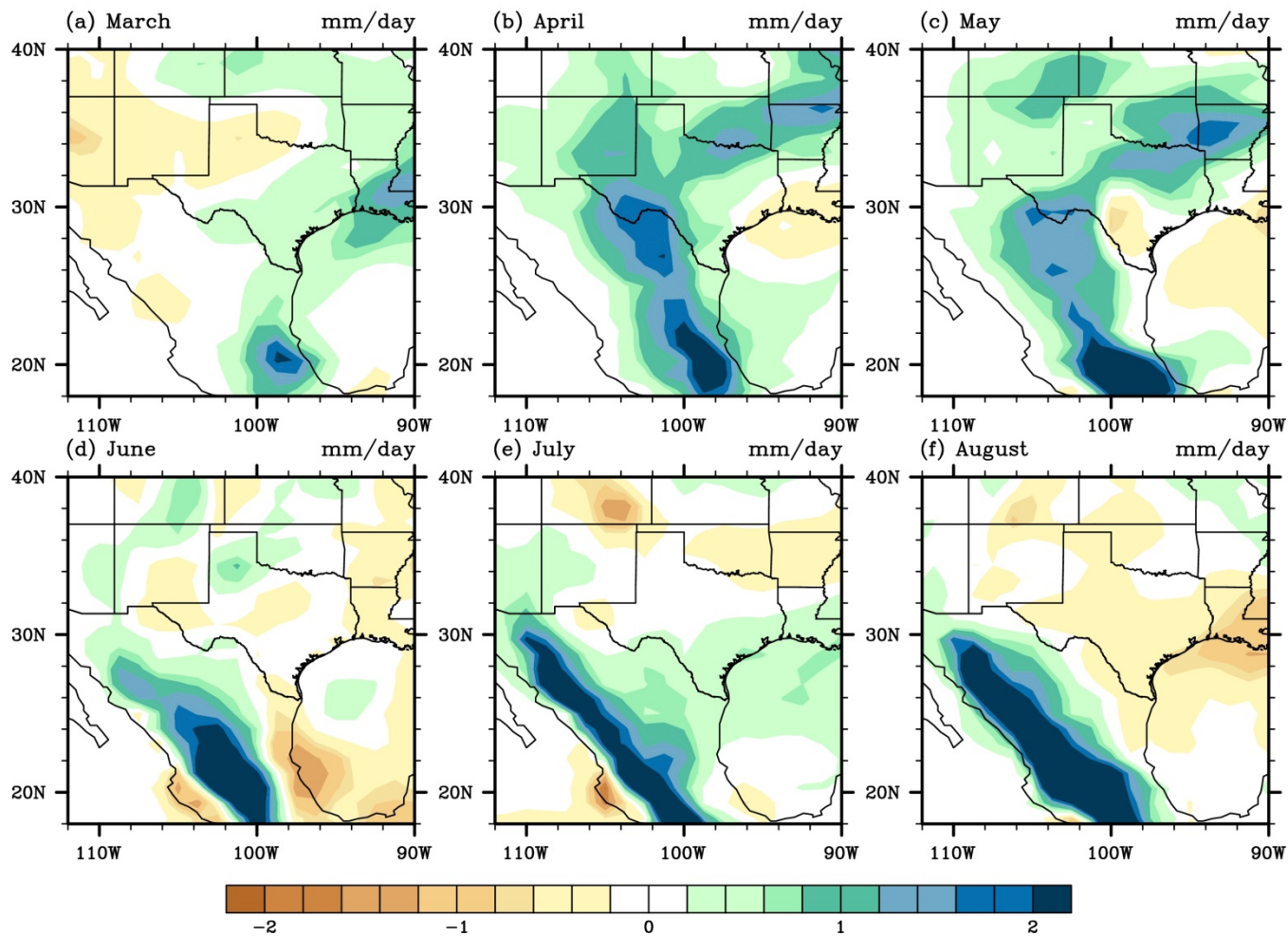
Rainfall decreases locally and downstream

E1(DRY) minus CTRL



Rainfall increases over the highlands and slightly decreases over Texas during summer

E2(WARM) minus CTRL



Conclusions

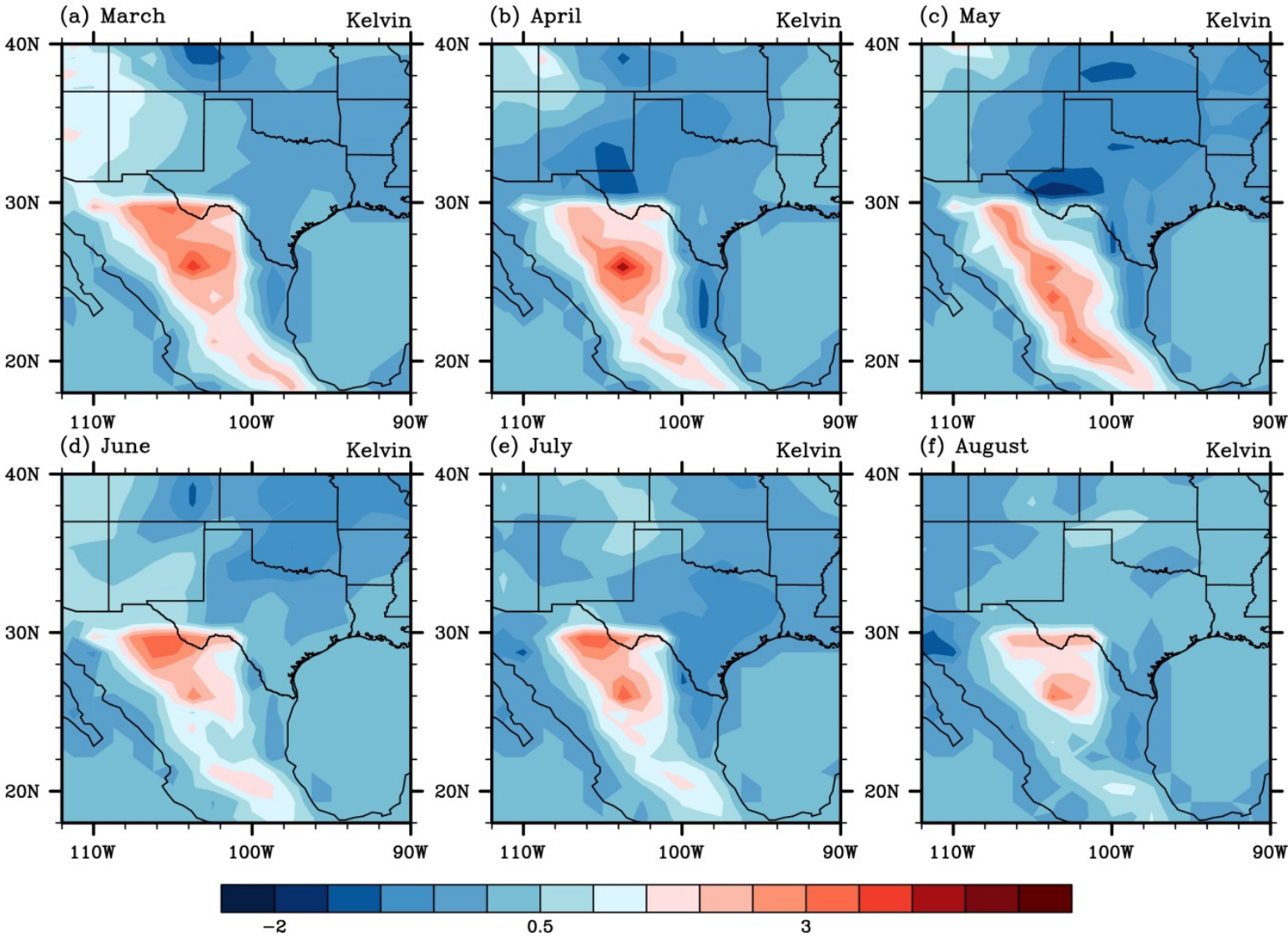
- *When the MP becomes dry, rainfall declines locally and downstream. During the spring, the dry air brought to Texas by prevailing westerly winds suppresses local convection; but dry air advection from the highlands has little influence on rainfall over Texas during the summer when Texas is no longer in the downstream areas.*
- *During the summer, a warmer MP acts like a “moisture pump” that pushes moist air over the peripheral low elevation areas to the highlands; it bends the low-level jet towards the highlands and an anti-cyclonic flow anomaly forms over the southern US, which causes air to diverge and tends to reduce rainfall over the southern US.*

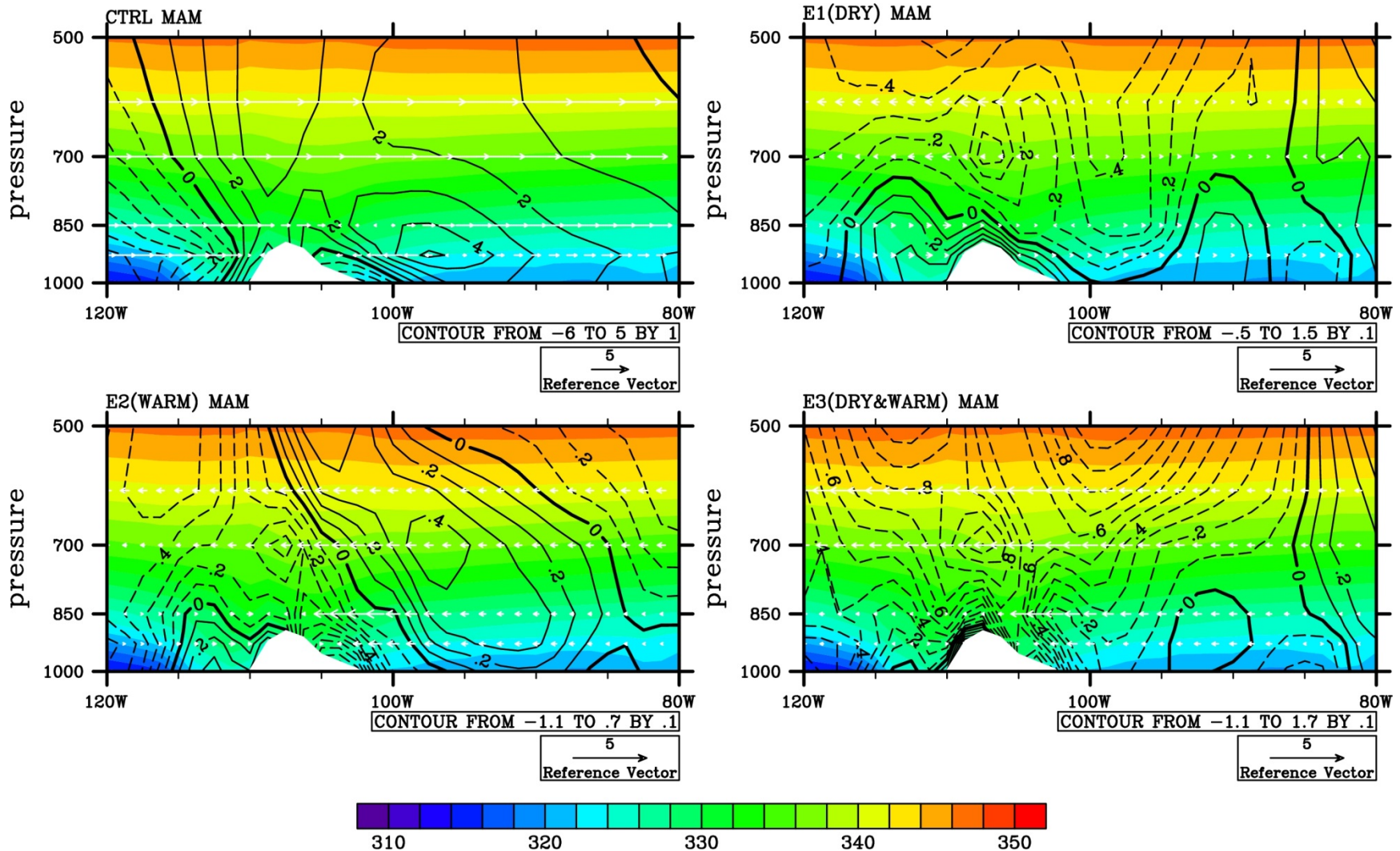


Thank you!

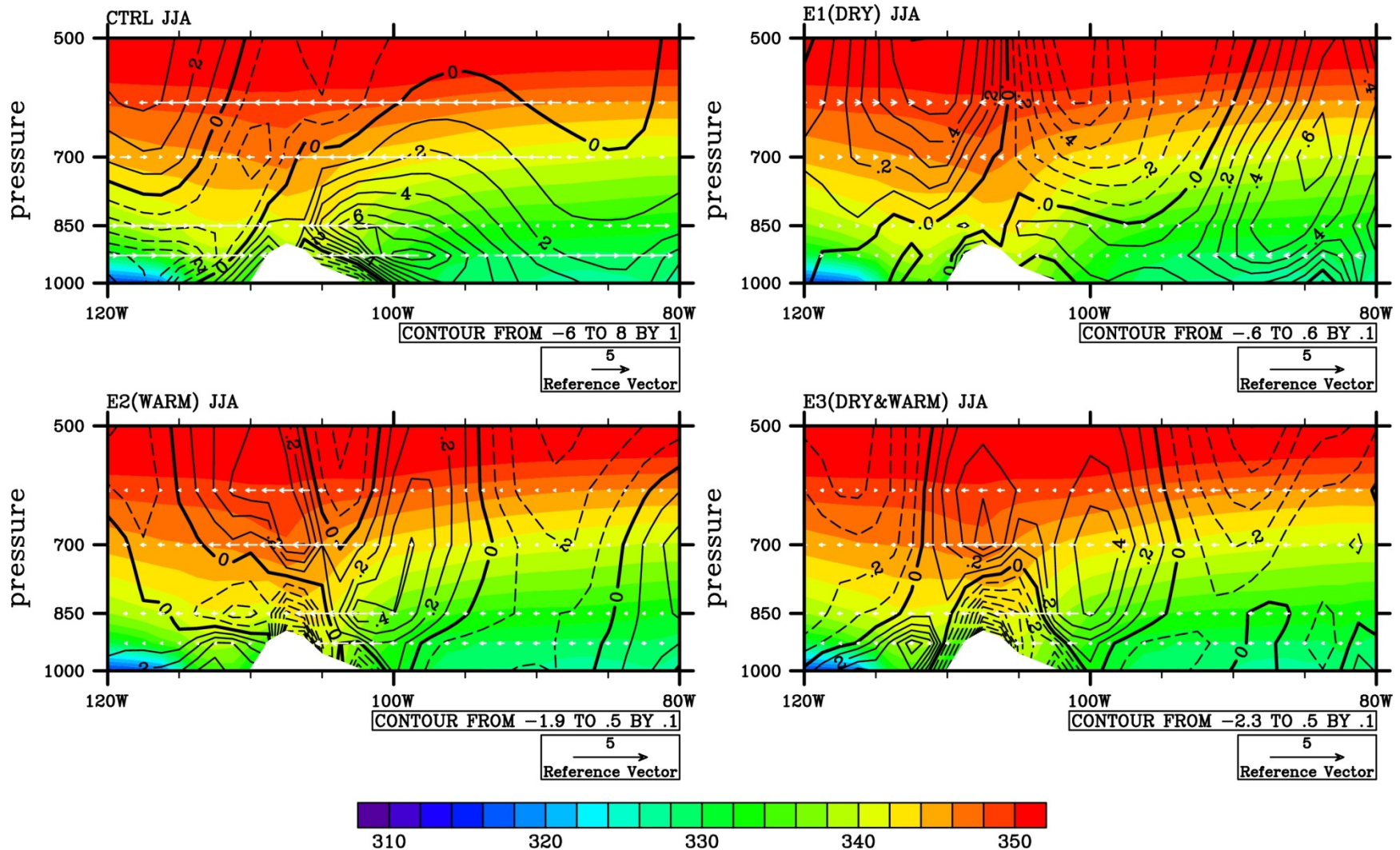
Surface temperature difference between E2 and CTRL

E2(WARM) minus CTRL





Vertical cross sections of potential temperature (K, shadings), meridional wind (m s⁻¹, contours), and zonal wind (m s⁻¹, white arrows) along 30°N from 120°W to 80°W for spring (MAM). The contours show meridional wind speed for CTRL and differences in meridional wind speed between the experiment runs and the control run for E1, E2, and E3.



Vertical cross sections of potential temperature (K, shadings), meridional wind (m s^{-1} , contours), and zonal wind (m s^{-1} , white arrows) along 30°N from 120°W to 80°W for spring (JJA). The contours show meridional wind speed for CTRL and differences in meridional wind speed between the experiment runs and the control run for E1, E2, and E3.