Growing water awareness



















Inclusion of a simple reservoir model in regionalscale surface water **Biver flow in the river basins of California**



modeling

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Integrated River Modeling



Atmospheric Model or Dataset



Land Surface Model



"Blue Line" River Network -High-Performance Computing River Network Model

RAPID

- Routing Application for Parallel computation of Discharge)
- Computes flow and optimizes model parameters
- Model code, input data and animations are available online
- Can run on supercomputers







http://www.ucchm.org/david/rapid.htm

Who made RAPID possible?

NCAR

MINES PARIS

UCAR Advance Study Program

CUAHSI

- CUAHSI Hydrologic Information System (NSF EAR-0413265)
- NASA Interdisciplinary Science Projects (NNX07AL79G and NNX11AE42G)
- Ecole des Mines de Paris
- AGU Horton (Hydrology) Research Grant
- David Maidment, Florence Habets, Zong-Liang Yang, David Gochis, Jay Famiglietti











Last we talked...





Water Forum 1 13 Feb 2012

David et al. (2013a Env. Model. & Soft)

Since then...



David et al. (2013b Water Resour. Res.)



Flow wave propagation and parallel computing



Also, moved to California





Reservoirs and river modeling



What does a reservoir do?



Daily flow for the year 2000, Shasta Lake

Treatment for reservoirs is needed!!!





Research motivations

- Add a simple reservoir model to RAPID that is based on natural physics
 - Advantages:
 - First step to reservoir modeling
 - Get started on coupling rivers/reservoirs
 - Limitations:
 - Does not explicitly account for storage of water as water supply and for corresponding management practices

Bernoulli Principle applied to reservoirs

 $\frac{v^2}{2} + g \cdot z + \frac{P}{\rho} = \text{constant}$

For incompressible and inviscid fluids



$$Q^{out} = \frac{1}{\tau} \cdot V \qquad \tau = \frac{A}{a} \sqrt{\frac{h}{2 \cdot g}}$$



1738

Daniel Bernoulli (1700-1782)

Assuming water height is almost constant

Linear reservoir equation

Response of a linear reservoir

 $Q^{out} = \frac{1}{\tau} \cdot V$

Linear reservoir equation

$$\frac{dV(t)}{dt} = Q^{in}(t) - Q^{out}(t)$$

Continuity equation (conservation of mass)

$$\frac{dQ^{out}(t)}{dt} + \frac{1}{\tau} \cdot Q^{out}(t) = \frac{1}{\tau} \cdot Q^{in}(t)$$

First order linear differential equation

Storage can be removed from the equation!!!

First order low-pass filter



Some review of existing literature

- Sherman, 1932, Streamflow from rainfall by the unit hydrograph method, Eng. New-Record
- Clark, 1945, Storage and the unit hydrograph
- Zoch, 1934, 1936, 1937, On the relation between rainfall and streamflow, *Monthly Weather Rev*.
- Nash, 1955, The relation of streamflow to rainfall, *ME Thesis, Univ College, Galway*
- Nash and Farrel, 1955, A graphical solution of the flood-routing equation for linear storage-discharge relation, *Trans. AGU*
- Dooge, 1956, Synthetic unit hydrographs based on triangular inflow, MS thesis, Univ. Iowa
- Nash, 1957, The form of the instantaneous unit hydrograph, *Int. Assoc. Hydrol. Sci. Gen. Assemb.*
- Nash, 1959, Systematic Determination of unit hydrograph parameters, *JGR*
- Nash, 1959, A note on the Muskingum flood routing method, JGR
- Dooge, 1959, A General Theory of the Unit Hydrograph, JGR



James Nash (1927-2000)



James Dooge (1922 – 2010)

Classic responses of low-pass filters



Unit impulse (Dirac delta function)





Square signal (or succession of unit step functions)



What if we assume first half of 16 sine wave ???

Response of linear reservoir to sine wave (analytical solution)

$$\left(\frac{dQ^{out}(t)}{dt} + \frac{1}{\tau} \cdot Q^{out}(t) = \frac{1}{\tau} \cdot Q^{in}(t)\right)$$

 $Q^{in}(t) = \sin\left(\frac{2\pi}{T} \cdot t\right)$

 $Q^{out}(0) = 0$

Differential equation

Inflow

Initial condition

$$\begin{cases} \mathcal{Q}^{out}\left(t\right) = \frac{1}{1 + \left(\frac{2\pi\tau}{T}\right)^{2}} \cdot \left[\frac{2\pi\tau}{T} \cdot \exp\left(\frac{-1}{\tau} \cdot t\right) + \sqrt{1 + \left(\frac{2\pi\tau}{T}\right)^{2}} \cdot \sin\left(\frac{2\pi}{T} \cdot \left(t - t_{\varphi}\right)\right) \right] \\ t_{\varphi} = \frac{T}{2\pi} \cdot \arctan\left(\frac{2\pi\tau}{T}\right) & \text{Exact solution} \end{cases}$$

$$17$$



T=20 τ=1



T=20 τ=2



T=20 τ=5



T=20 τ=10



T=20 τ=20



T=20 τ=40



T=20 τ=100

What did we observe???



T=20 τ=10

Some characteristics of these hydrographs

- The outflow seems to be a periodic signal of same period as inflow, but:
 - with a different phase (outflow delayed)
 - with a different magnitude (outflow smaller)
- The difference in phase seems to increase with the value of τ/T but reaches a maximum for large values of τ /T
- The difference in magnitude seems to increase with the value of τ/T
- The initial peak of the outflow is higher than the following peaks

Quantification of the difference in phase

$$t_{\varphi} = \frac{T}{2\pi} \cdot \arctan\left(\frac{2\pi\tau}{T}\right)$$

$$\begin{aligned} T &> 0 \\ \tau &> 0 \end{aligned} t_{\varphi} \in \left] 0, T / 4 \right[\end{aligned}$$

$$\frac{\tau}{T} \square 1 \Longrightarrow t_{\varphi} \square 0, t_{\varphi} > 0$$
$$\frac{\tau}{T} \square 1 \Longrightarrow t_{\varphi} \square \frac{T}{4}, t_{\varphi} < \frac{T}{4}$$

$$\tau = T \cdot \frac{1}{2\pi} \cdot \tan\left(\frac{2\pi}{T} \cdot t_{lag}\right)$$

If t_{lag} is close to T/4, (i.e. $\tau/T >>1$) impossible to get an accurate estimate of τ

Some maths...

$$\max_{t \in [0,T]} \left[Q^{out}(t) \right] = \frac{1}{1 + \left(\frac{2\pi\tau}{T}\right)^2} \cdot \left[\frac{2\pi\tau}{T} \cdot \exp\left(\frac{-T}{2\tau}\right) + \sqrt{1 + \left(\frac{2\pi\tau}{T}\right)^2} \right]$$



Taylor's series expansion

Brook Taylor (1685-1731)

- Resolution of a second order polynomial
- Assumption that max[Q^{out}] is small

$$\frac{\tau}{T} = \frac{1}{\pi \cdot \max_{t \in [0,T]} \left[Q^{out} \left(t \right) \right]} - \frac{1}{4}$$

Valid for t_{lag} close to T/4 (i.e. $\tau/T>>1$)

Application to Shasta Lake



Jan – Mar 2000, daily

Rising limb of Qⁱⁿ=5 days Rising limb of Q^{out}=5 days T_{lag} =5 days Max[Qⁱⁿ]=73,332 cfs Max[Q^{out}]=40,776 cfs

$$\tau = 4 \cdot 5 \cdot \frac{1}{\pi \cdot \frac{44776}{73332}} - \frac{1}{4} = 6.45 days$$

Numerical approximation for linear reservoir

$$\frac{dQ^{out}(t)}{dt} = \frac{Q^{out}(t + \Delta t) - Q^{out}(t)}{\Delta t}$$

$$Q^{out} = \frac{1}{\tau} \cdot V$$

First order explicit Euler Method applied to Continuity Eq

Linear reservoir Eq



Leonhard Euler (1707-1783)

$$Q^{out}\left(t+\Delta t\right) = \frac{\Delta t}{\tau} \cdot Q^{in}\left(t\right) + \left(1 - \frac{\Delta t}{\tau}\right) \cdot Q^{out}\left(t\right)$$

Numerical method for linear reservoir



T=20 τ=1



T=20 τ=2



T=20 т=5



T=20 τ=10



T=20 τ=20



T=20 τ=40



T=20 τ=100

Reservoir modeling of Shasta Lake



Jan – Mar 2000, daily

RAPID with and without linear reservoirs



Applied to the two biggest reservoirs in California.

RMSE and efficiency improved a the five corresponding downstream stations.

Results are statistically significant!!!

Shasta Lake

Hydrological Sciences - Journal - des Sciences Hydrologiques, 35, 6, 12/1990

The education of hydrologists (Report of an IAHS/UNESCO Panel on hydrological education)

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Thank you!

It is fashionable to regard man's involvement in nature as almost always bad. It is true that ignorance (often spurred on by greed) is leading to progressive damage to the natural environment and that we require as much scientific understanding as possible of the relevant processes in order to diagnose our mistakes and put them right. On a more positive note, however, man's ability to control his environment, to improve it and to make it more enjoyable, and indeed more productive and profitable, depends just as centrally on putting our understanding of hydrological processes on as sound a scientific basis as we can manage.