

## Chlorine and hydrogen isotope geochemistry of obsidian glasses: behavior during volcanic degassing at Mono Craters, CA

Prather, T.<sup>1</sup>, Barnes, J.D.<sup>1</sup>, Befus, K.<sup>1</sup>, Cisneros, M.<sup>1</sup>, and Gardner, J.E.<sup>1</sup>

[timothy.h.prather@utexas.edu](mailto:timothy.h.prather@utexas.edu)

*1. Jackson School of Geosciences, University of Texas, Austin, Texas 78712, USA*

Volatile element concentrations (Cl, H<sub>2</sub>O, and CO<sub>2</sub>) and stable isotope compositions ( $\delta$ D and  $\delta^{37}\text{Cl}$ ) of volcanic glasses (obsidians) ( $n = 30$ ) have been determined to quantify the behavior of chlorine stable isotopes ( $^{35}\text{Cl}$  and  $^{37}\text{Cl}$ ) during volcanic degassing. Pyroclastic obsidian samples were from tuff layers representing a single eruptive sequence that occurred around 1350 A.D. in the Mono Craters volcanic field, California. The Cl, H<sub>2</sub>O, and CO<sub>2</sub> concentrations recorded by these eruptive obsidians track the chemical evolution of the magmatic system. The H<sub>2</sub>O and CO<sub>2</sub> concentrations of the samples range from 0.37 to 2.08 wt% and 1 to 31 ppm, respectively. H<sub>2</sub>O and CO<sub>2</sub> concentrations are strongly correlated and reflect the degassing trend of the eruptive sequence. Chlorine concentrations of obsidians range from 609 to 833 ppm and do not display a strong correlation with either H<sub>2</sub>O or CO<sub>2</sub> concentrations. Obsidians were selected from two tuff layers: 1) a lower layer containing average H<sub>2</sub>O and CO<sub>2</sub> concentrations of  $1.5 \pm 0.5$  wt% and  $22 \pm 11$  ppm, respectively, and 2) an upper layer containing slightly lower average H<sub>2</sub>O and CO<sub>2</sub> concentrations of  $0.9 \pm 0.5$  wt% and  $6 \pm 5$  ppm, respectively. Chlorine concentrations are essentially identical between the two layers, averaging  $742 \pm 58$  ppm in the lower layer and  $702 \pm 75$  ppm in the upper layer.

Measured  $\delta$ D values of the obsidians vary between -63 to -74‰ ( $1\sigma = \pm 2\%$ ) and display D/H ratios that decrease with lower total water content following a distillation trend dominated by open system degassing.  $\delta^{37}\text{Cl}$  values were measured on select samples from each of the two tuff layers. The samples from the lower layer have  $\delta^{37}\text{Cl}$  values between -1.8 to -2.0‰ ( $n = 3$ ), whereas the samples from the upper layer have  $\delta^{37}\text{Cl}$  values between -1.3 and -1.4‰ ( $n = 3$ ) ( $1\sigma = \pm 0.2\%$ ). Despite the similar Cl concentrations between the two layers, the samples with lower  $\delta^{37}\text{Cl}$  values have higher Cl concentrations ( $763 \pm 61$  ppm Cl) than samples with higher  $\delta^{37}\text{Cl}$  values ( $639 \pm 33$  ppm Cl). The samples analyzed for Cl isotopes span the range of observed H<sub>2</sub>O and CO<sub>2</sub> concentrations: a sample with one of the highest H<sub>2</sub>O and CO<sub>2</sub> concentrations (2.1 wt% and 26 ppm, respectively) has a  $\delta^{37}\text{Cl}$  value of -1.8‰; whereas, a sample with one of the lowest H<sub>2</sub>O and CO<sub>2</sub> concentrations (0.4 wt% and 3 ppm, respectively) has a  $\delta^{37}\text{Cl}$  value of -1.4‰. Preliminary data indicates the possibility of minor chlorine isotope fractionation during the eruptive sequence at Mono Craters; however, additional work will better quantify the magnitude of this fractionation.

**Keywords:** chlorine isotopes, degassing, Mono Craters, obsidian